

Towards a World Height System – a Status Report

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On the Basis of a technical paper prepared together with
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- **Inter-Commission Project 1.2: Vertical Reference Frames (Conventions for the definition of World Height System, 2003 – 2011)**
- **GGOS Theme 1: Unified Global Height System (Action Items for the unification Height Reference Systems, 2011 until now)**
- **ESA project “GOCE+ Height System Unification with GOCE”**
- **Inventory of standards and conventions used for the generation of IAG/GGOS products by GGOS BSC**
- **Joint Working Group 0.1.1: Vertical Datum Standardisation**
- **In 2014 a position paper “Proposal for the Definition and Realization of an International Height Reference System (IHRIS)” was prepared, distributed in Nov. 2014.**
- **The IAG Executive Committee decided in December 2014, to establish an Ad-hoc Group on an International Height Reference System (IHRIS)**

- **Motivation for a Global Height Reference System**
- **Height and Earth Gravity Field**
- **Concepts for Unification of Height Reference Systems**
- **Numerical standards**
- **Tidal Systems**
- **Definition of an International Height Reference System**
- **Realization of an IHRS**
- **Integrated Networks**
- **Next Steps**

Motivation for a Global Height Reference System

External view

- Climate change is one of the themes debated in politics today and one of the general developments which influence the living conditions of mankind
- Sea level rise is one of the mayor risks for urban regions in many parts of the world
- A common global vertical reference system is the metrological basis for monitoring effects related to the vertical scale (geometry and Earth gravity field)

Geodetic view

- Inconsistencies between geometric products and products related to the Earth gravity field need to be removed to enable the development of integrated geodetic applications.
- Determination of time-dependent changes of the vertical reference frame needs long term consistent stable reference.

Physical Heights and Earth Gravity Field

The representation of the Earth gravity field is independent possible with to different kind of fields.

Both fields

- geo-potential scalar field $W(X)$
 - the outer Earth gravity vector field $\vec{g}(X)$
- connected by the relationship

$$\vec{g}_P = \text{grad } W_P = -g_P \begin{pmatrix} \cos \Phi & \cos \Lambda \\ \cos \Phi & \sin \Lambda \\ & \sin \Phi \end{pmatrix}, \quad g_P = g(X) = |\text{grad } W_P|$$

X position in natural coordinates: Φ astronomical latitude,
 Λ astronomical longitude, W potential of Earth gravity field.

Concepts for unification of height reference systems

General possibilities for the unification

- 1) On continents:
by common adjustment of existing levelling networks (cPk)
- 2) Global:
general case for realization and unification - combination by global GGM and GNSS/levelling
- 3) Global Over oceans:
using a model of mean sea surface topography and tide gauge observations

Combinations are useful and some times necessary

Normals Amsterdams Pail

an early warning system, to warn and to protect the citizens of Amsterdam of floods.

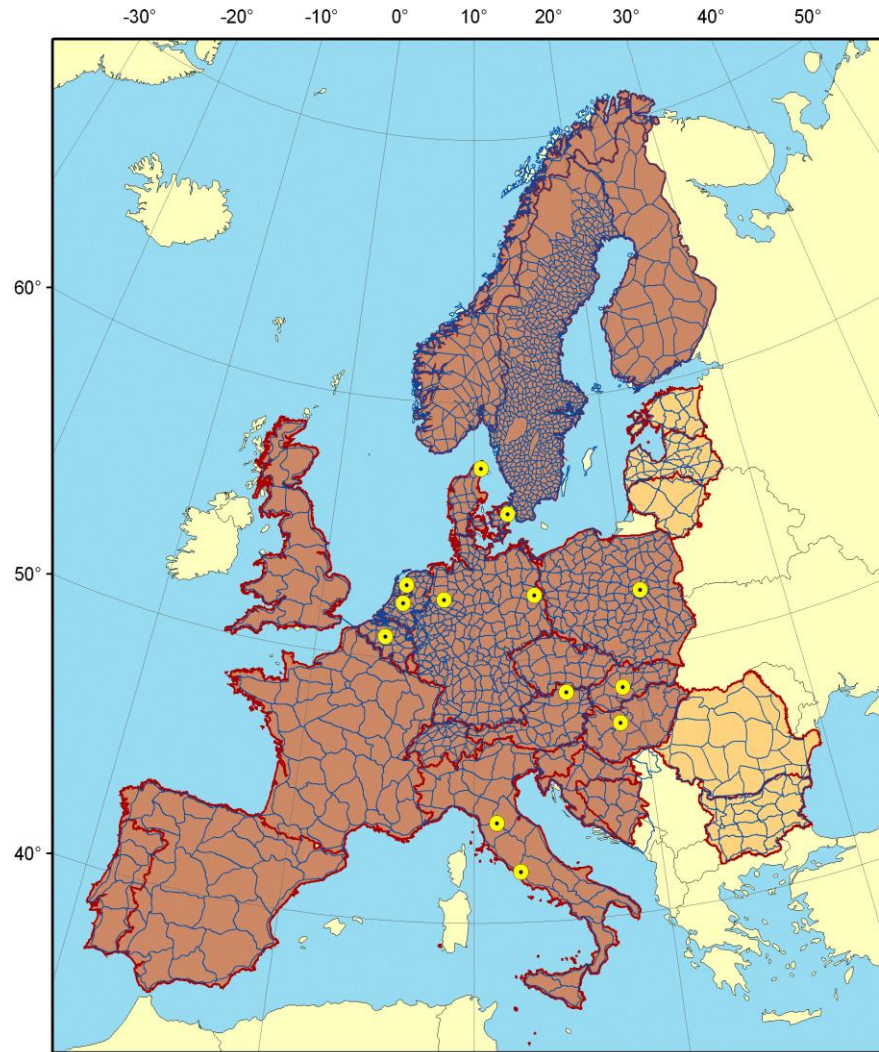
Mathematician and Mayor Johnnes Hudde promoted the measurement of water levels at high and low tide from September 1, 1683 to September 1, 1684

At 8 gates of the new water weir height marks were attached ("dijkpeilstenen" = stones that mark the height of the dike):
Inscription: Seedeichhöhe equal to nine feet five inches (= 2,6789m) on city level.



Last original „dykpeilsteen“ in der „Nieuwe Burg“ gate

1) EVRS: Common adjustments of existing levelling networks



- Realization of EVRF2000 datum: reference point *000A2530 in the Netherlands*
- Realization of EVRF2007 datum
 - several datum points distributed over the stable part of Europe
 - participating countries were asked to propose stable points
 - 13/19 points have been used

up to 1998
as from 2003
UELN lines

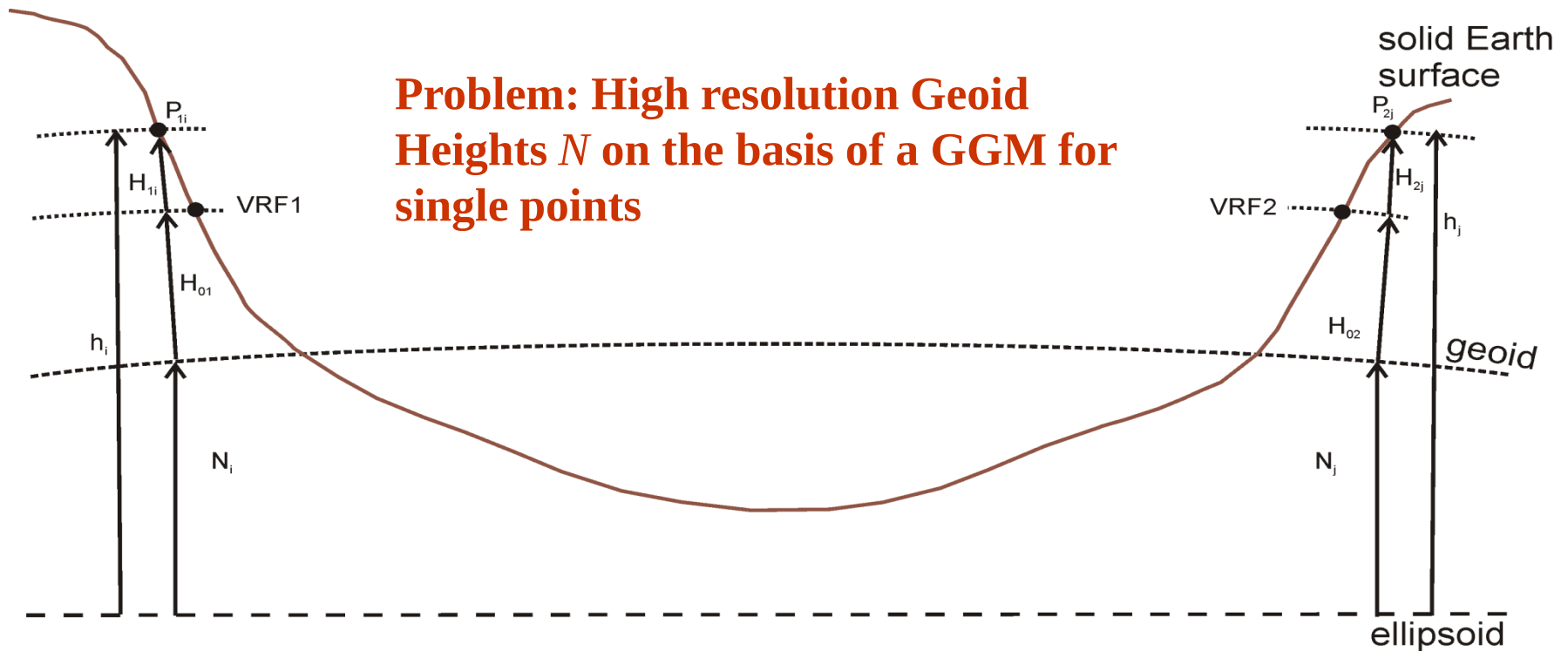
Unification and realization by global GGM and GNSS/levelling

$$H_{0,VRF} = h_{i,ITRF} - H_{i,VRF} - N_{i,GGM}, \text{ } h \text{ and } N \text{ shall global}$$

VRF realization in single points:

$$H_{VRF} = h_{ITRF} - N_{GGM + local}$$

Problem: High resolution Geoid Heights N on the basis of a GGM for single points



Numerical Standards

The GRS80

For all geodetic tasks we need standards, conventions, guidelines, rules!

Geodetic Reference System 1980 (GRS 80) defines major parameters for geodetic reference systems related to a level ellipsoid

- **Agreed by IUGG, IAG and IAU in 1979.**
- **Recommended by IAG for the conversion of ITRF Cartesian coordinates to ellipsoidal coordinates.**
- **Worldwide use for many map projections and million of coordinates, e.g. ETRS89**

Numerical Standards

IUGG GA 1991

- **At the IUGG General Assembly 1991 in Vienna new values for the geocentric gravitational constant GM and the semi-major axis a of the level ellipsoid were recommended.**
- **Since this time these parameters have been used in global gravity models e.g. EGM96**



Numerical Standards Conflict

- In IERS 2003 and 2010 conventions are two sets of parameters of a level ellipsoid in use: In numerical standards and the GRS 80
- The gravitational constants GM of GRS80 and IERS 2010 conventions differ by about $0.9 \text{ m}^3\text{s}^{-2}$, and the semi-major axis of both standards differs by 0.4 m
- Also noteworthy is that the IERS 2010 conventions recommend different level ellipsoid parameters for different applications.
- For the reduction of Cartesian coordinates into ellipsoidal ones, the GRS80 Ellipsoid is recommended for application

Numerical Standards

Sets of conventional parameters

ellipsoid	Semi-major axis a in m	flattening f^{-1}	Geocentric gravitational constant GM in $10^8 \text{m}^3 \text{s}^{-2}$	U_0/W_0 in $\text{m}^2 \cdot \text{s}^{-2}$	Y_e in $\text{m} \cdot \text{s}^{-2}$
Int. Ell. 1930 (Hayford)	6 378 388	297	3 986 329		
GRS 67	6 378 160	298.247	3 986 030		
GRS 80	6 378 137	298.257222101	3 986 005	6 263 6860.850	9.78032 677
WGS 84	6 378 137	298.25722356			
IUGG 91	6 378 136.3 0.5		3 986 004.41 0.01		
IERS 2003 Conventions (zero tide)	6 378 136.6 0.1	298.25642 0.00001	3 986 004.418 0.008	6 263 6856.0 0.5	(9.78032 666)
EGM96	6 378 136.3		3 986 004.415		
EIGEN CG01C (tide free)	6 378 136.46		3 986 004.415		

Angular velocity of the Earth rotation **7 292 115** $10^{-11} \text{ rad s}^{-1}$

IAG has two complete sets of parameters in use

Numerical Standards

Defining parameters of level ellipsoids (equipotential or mean Earth ellipsoid)

Ellipsoid	Semi-major axis a in m	Flattening f^{-1}	Geocentric gravitational constant GM in $10^8 \text{m}^3 \text{s}^{-2}$	Normal/geoidal potential U_0/W_0 in $\text{m}^2 \cdot \text{s}^{-2}$	Normal gravity at equator γ_e in $\text{m} \cdot \text{s}^{-2}$
GRS 80	6 378 137	298.2572221 ± 01	3 986 005	6 263 6860.850	9.78032677
IERS Conventions (2010) (zero tide)	6378136.6 ± 0.1	298.25642 ± 0.00001	3986004.418 ± 0.008	62636856.0 ± 0.5	
GRS20XX	tbd.	298.25642 ± 0.00001	3986004.418 ± 0.008	tbd.	tbd.

Angular velocity of the
Earth rotation ω
in

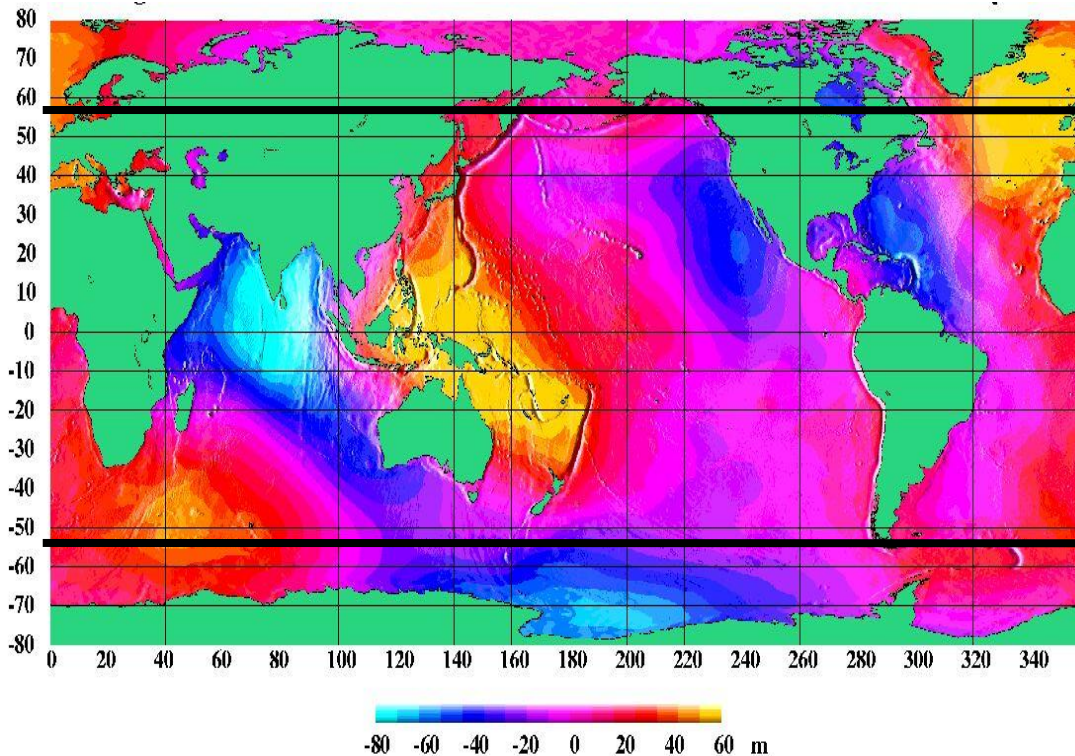
7 292 115
 $10^{-11} \text{rad s}^{-1}$

Determination of W_0 of Mean Sea Surface

$$W_S = U_0 + \partial U_0 / \partial h \cdot h_{SALT} + T_{SGGM}$$

$$W_{0S} = 1/S \iint_S W_S dS$$

- With satellite altimeter observations (ALT) and a global gravity model (GGM)
- in an agreed area of free oceans
 - over a defined time period
 - at a defined epoch.



KMS04 MSS model

Should the W_0 value be changed?



**The consequences
have to be considered!**

Carl Calvert, EUREF 2007

Carl
8.6.7

The current best-estimated value for W_0 shall be defined (and frozen) as the potential value of the geoid and reference for a global height reference system. The procedure of W_0 determination has to be documented in conventions/guidelines.

Numerical Standards

Tidal Systems - Reduction of data

gravity	geoid	levelling height	altimetry	mean sea level	position
$g/\Delta g$	W/N	ΔH	h	msl	X/h

Mean tidal system

Mean/zero crust

(Stokes is not valid if masses outside the Earth surface)

Δg_m

N_m

ΔH_m

Relation to N_m for
oceanographic
studies h_{msl}

Zero tidal system

Mean/zero crust

(Recommended by IAG
Res. No. 16, 1983)

Δg_z Stokes N_z
→ (EGG97)

ΔH_z
 c_p

Tide-free system

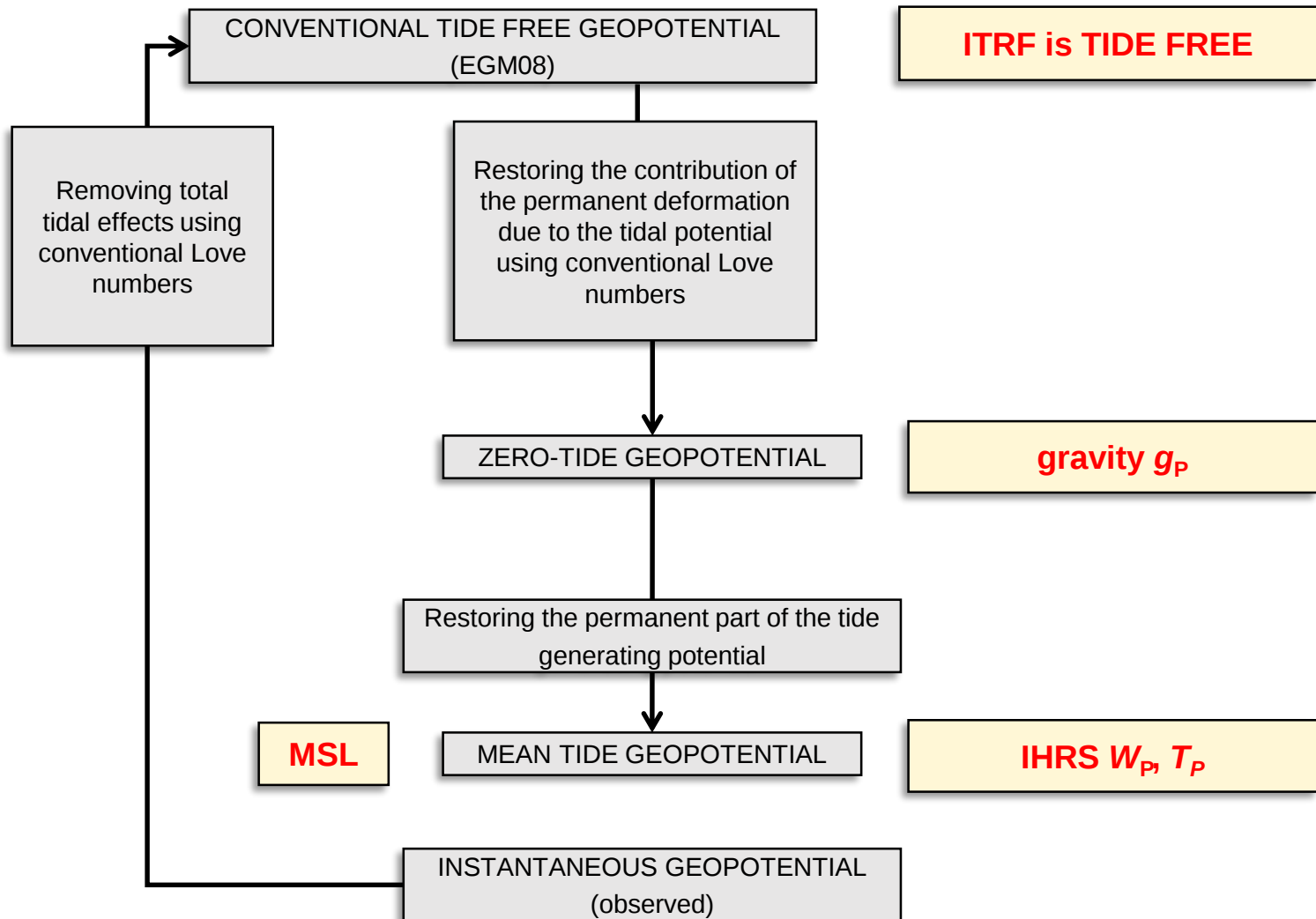
Tide-free crust

(unobservable, far away from
the real earth shape – there is
no reason for the non tide)

Δg_n Stokes N_n
→ (EGM96)

X_n
ITRFxx,
ETRS89

Tidal Systems and Reference Systems



Definition, realization and unification of height reference systems

- **Following the praxis of to handle geodetic reference system concepts we have to consider the definition and the realization of a reference system**
- **For this we need conventions (standards) and guidelines**
- **Unification of height systems is part of the realization**



Definition of an International Height Reference System (IHRs)

Five conventions define the IHRs (Status May 2015):

- The vertical reference level is the normal potential (or geopotential at the geoid or the geoid potential parameter) W_0 as an equipotential surface of the Earth gravity field. The relationship between W_0 and the Earth body must be defined and reproducible. $U_0 = W_0$ as a defining parameter of the conventional geocentric level ellipsoid is under discussion.
- Parameters, observations, and data shall be related to the mean tidal system/mean crust. (Under discussion)
- The unit of length is the meter (SI). The unit of time is the second (SI). This scale is consistent with the TCG time coordinate for a geocentric local frame, in agreement with IAU and IUGG (1991) resolutions, and is obtained by appropriate relativistic modeling.

- The vertical coordinates are the differences $-\Delta W_P$ between the potential W_P of the Earth gravity field at the considered points P , and the geoidal potential of an conventional W_0 . The potential difference $-\Delta W_P$ is also designated as geopotential number c_P :

$$-\Delta W_P = c_P = W_0 - W_P$$

- The spatial reference of the position P for the potential $W_P = W(X)$ is related as coordinates X of the International Terrestrial Reference System.

Proposal for the elements of an IHRF:

- The **geopotential at the geoid W_0** is developed through best estimates. The procedure of W_0 determination must be documented in conventions/guidelines, to ensure the reproducibility and interpretability of changes.
- The central element of the IHRF is a **Global Gravity Model (GGM)**. One (combined) satellite only GGM for homogenous long wavelength approximation of the Earth gravity potential as conventional. One GGM combined with terrestrial data for direct application in sparsely surveyed regions.

- **A network of geodetic observation stations:** The network includes geodetic observatories, absolute gravity stations, key comparison stations of an IGSN, selected tide gauge stations (TIGA-IGS), time stations.

At the stations the

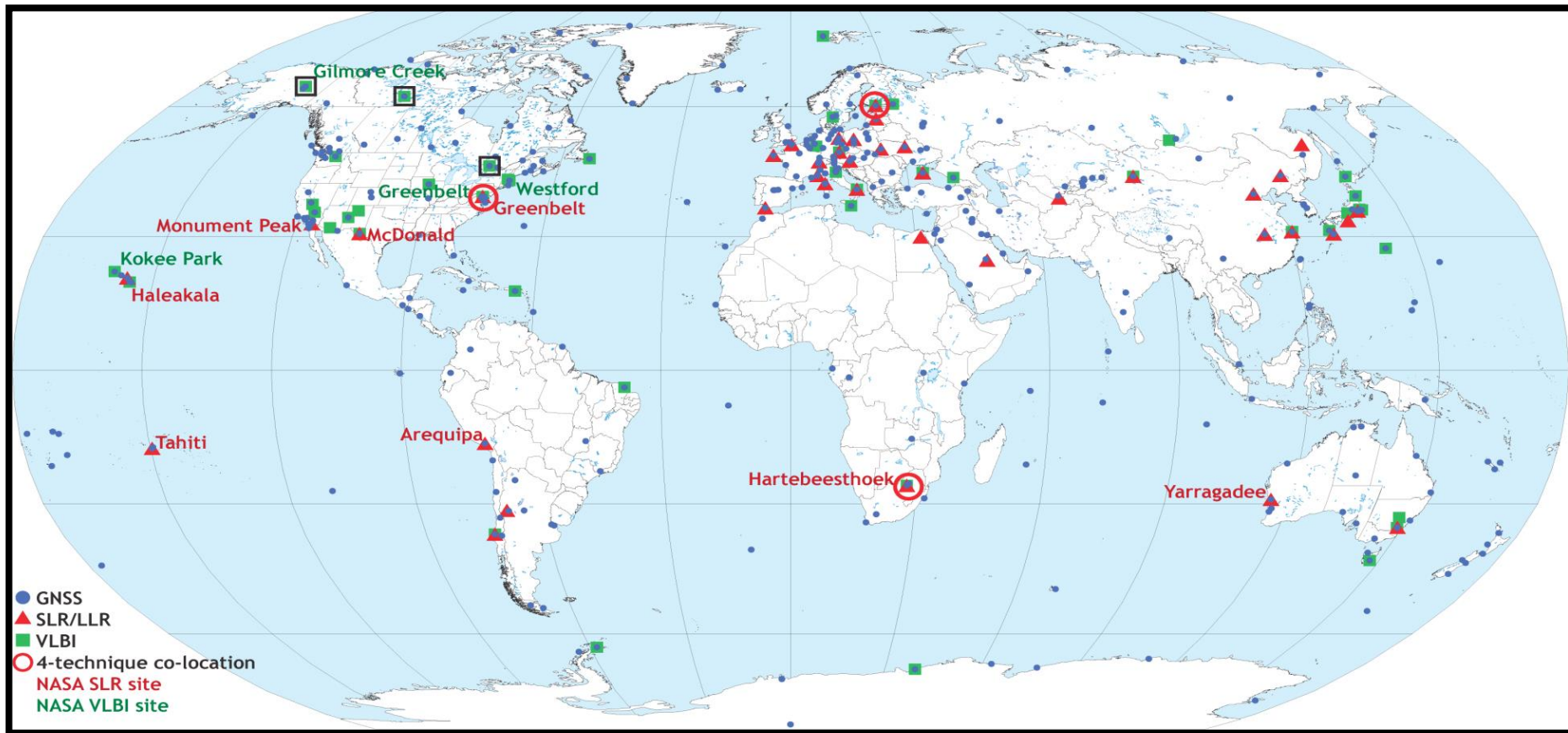
- Earth gravity potential W_p in IHRS
- Spatial reference X in ITRS
- Absolute gravity g

must be determined with highly possible accuracy.

This are integrated geodetic networks.

Integrated Networks

Global Network of Geodetic Stations and Observatories



- Common discussion of the members of the action group and the Ad-hoc Group about the open questions
- Editing of technical papers for the documentation of the conventions
- Preparation of relevant resolutions for the IUGG/IAG GA 2015

