

A combined geoid model for Hungary based on the spectral properties of gravity related data

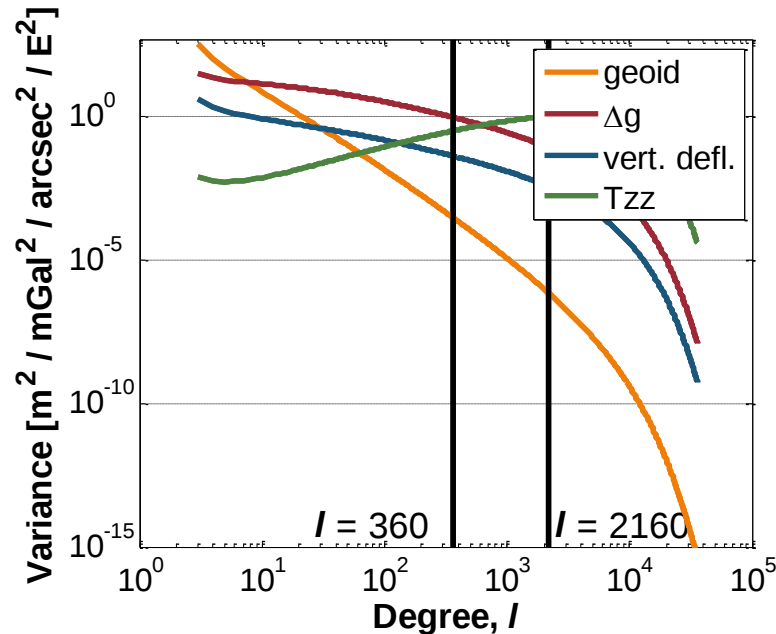
Eszter Szűcs

Geodetic and Geophysical Institute
Research Centre for Astronomy and Earth Sciences
Hungarian Academy of Sciences

EUREF2013 Symposium
29-31, May 2013 Budapest



Spectral sensitivity of gravity field quantities



Tscherning- Rapp model

$$\sigma_l^2(\Delta g) = \frac{A(l-1)}{(l-2)(l+B)} \sigma_0^{l+2}$$

$$A = 425.28 \text{ mGal}^2$$

$$B = 24$$

$$\sigma_0 = 0.999617$$

	3 - 36	37 - 360	361-2160	2161-36000
N	99.2 %	0.8 %	0.0 %	0.0 %
Δg	22.5 %	42 %	28.6 %	6.9 %
vertical defl.	29.5 %	38.6 %	25.6 %	6.3 %
T_{zz}	0.0 %	0.8 %	19.1 %	80.1 %

Spectral combination

$$T(r, \phi, \lambda) = T^M(r, \phi, \lambda) + T^G(r, \phi, \lambda) + T^E(r, \phi, \lambda)$$

GGM

EGM2008

gravity
anomaly

horizontal
grav. gradients

optimal combination in the spectral domain

n-th degree surface harmonics of T

$$T_n = w_n^M T_n^M + w_n^G T_n^G + w_n^E T_n^E \quad w_n^M + w_n^G + w_n^E = 1, \quad \forall n$$

$$T^{M,G,E} = \sum_{n=2}^{n_{max}} \left(\frac{a}{r}\right)^{n+1} w_n^{M,G,E} T_n^{M,G,E}$$

Spectral combination

T^M from GGM

$$T^M(r, \phi, \lambda) = \sum_{n=2}^{n_{max}} w_n^M T_n^M$$

T^G from gravity anomalies

$$T^G(r, \phi, \lambda) = \frac{R}{4\pi} \iint_{\sigma} S^*(\psi) \Delta g d\sigma \quad S^*(\psi) = \sum_{n=2}^{n_{max}} \frac{2n+1}{n-1} w_n^G P_n(\cos\psi)$$

T^E from the combination of T_{zx} and T_{zy} gradients

$$T^E(r, \phi, \lambda) = \frac{R^2}{4\pi} \iint_{\sigma} E^*(\psi) (T_{zy} \cos\alpha^* + T_{zx} \sin\alpha^*) d\sigma$$
$$E^*(\psi) = \sum_{n=2}^{n_{max}} \frac{1}{n(n+1)} w_n^E P_n^1(\cos\psi)$$

Spectral weights

weighting scheme

deterministic

not involve error estimates

stochastic

full error knowledge

e.g. EGG2008

quasi-deterministic (Kern et al.)

$$\tilde{\sigma}_n = \sigma_n + \varepsilon_n \approx \sigma_n^{model} + \varepsilon_n$$

from
data

"true"
unknown

σ_n / ε_n : amount of signal /error
contained in each degree

$$\varepsilon_n = \left| \tilde{\sigma}_n - \sigma_n^{model} \right|$$

data

GGM or
analytical model (Kaula, ...)

Kern et al. (2007).: A study on the combination of satellite, airborne, and terrestrial gravity data. J Geod. 77: 217-225.

Spectral weights – QD method

Error degree variances

GGM

$$\varepsilon_n^{GGM} = \sum_{m=1}^n (\varepsilon_{Cnm}^2 + \varepsilon_{Snm}^2)$$

terrestrial data

i-th measurement (*i* = 1, 2, ...)

$$\varepsilon_n^i = \left| \tilde{\sigma}_n^i - (\sigma_n^{model})^i \right|$$

Weights for $\forall n$

	$n \leq n_{\max}(\text{GGM})$	$n > n_{\max}(\text{GGM})$
GGM	$w \propto 1/\varepsilon_n^{GGM}$	$w = 0$
<i>i</i> -th data set	$w \propto 1/\left \tilde{\sigma}_n^i - (\sigma_n^{GGM})^i \right $	$w \propto 1/\left \tilde{\sigma}_n^i - (\sigma_n^{model})^i \right $

Degree variances from PSD

2D gridded data



2D PSD



1D radial
PSD

planar power spectrum, $PSD(w)$

$$\sigma_n = \frac{n + 0.5}{2\pi R^2} PSD(w) \quad w \approx \frac{n + 0.5}{R}$$



spherical power spectrum

degree
variances

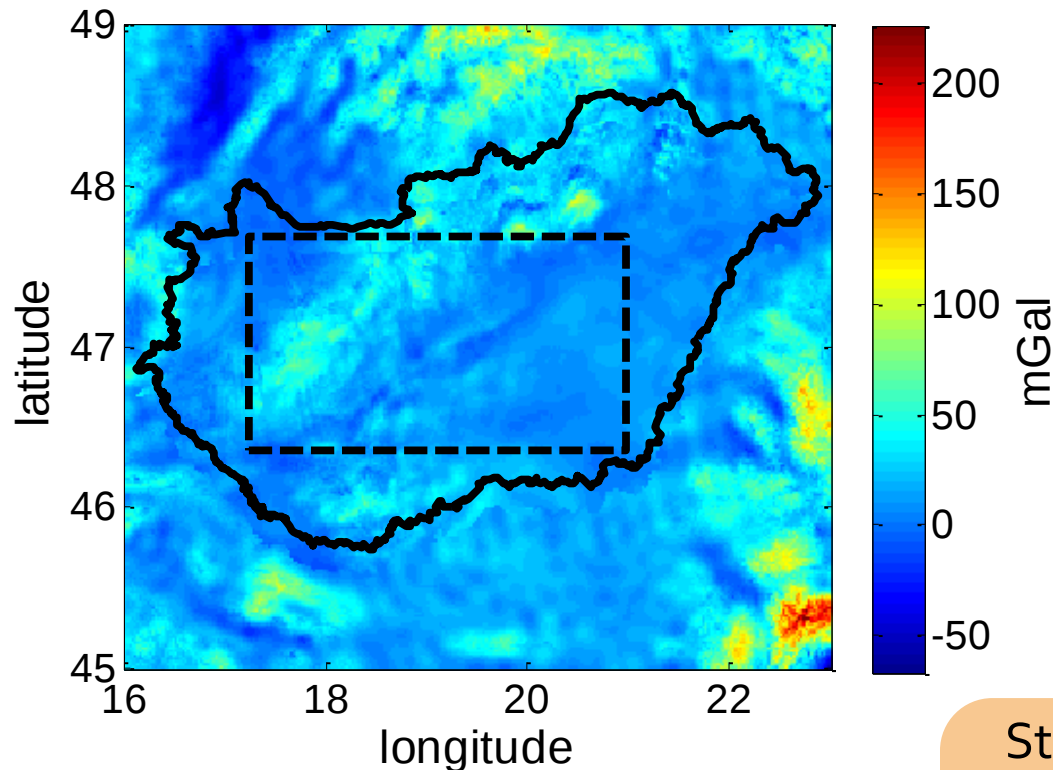
resolvable spectral band

$$n_{min} = \frac{40000 \text{ km}}{D \text{ [km]}} \quad \mathbf{D:} \text{ extension of the data set}$$

$$n_{max} = \frac{20000 \text{ km}}{d \text{ [km]}} \quad \mathbf{d:} \text{ grid spacing}$$

Data sets - Gravity

RTM reduced free-air anomaly



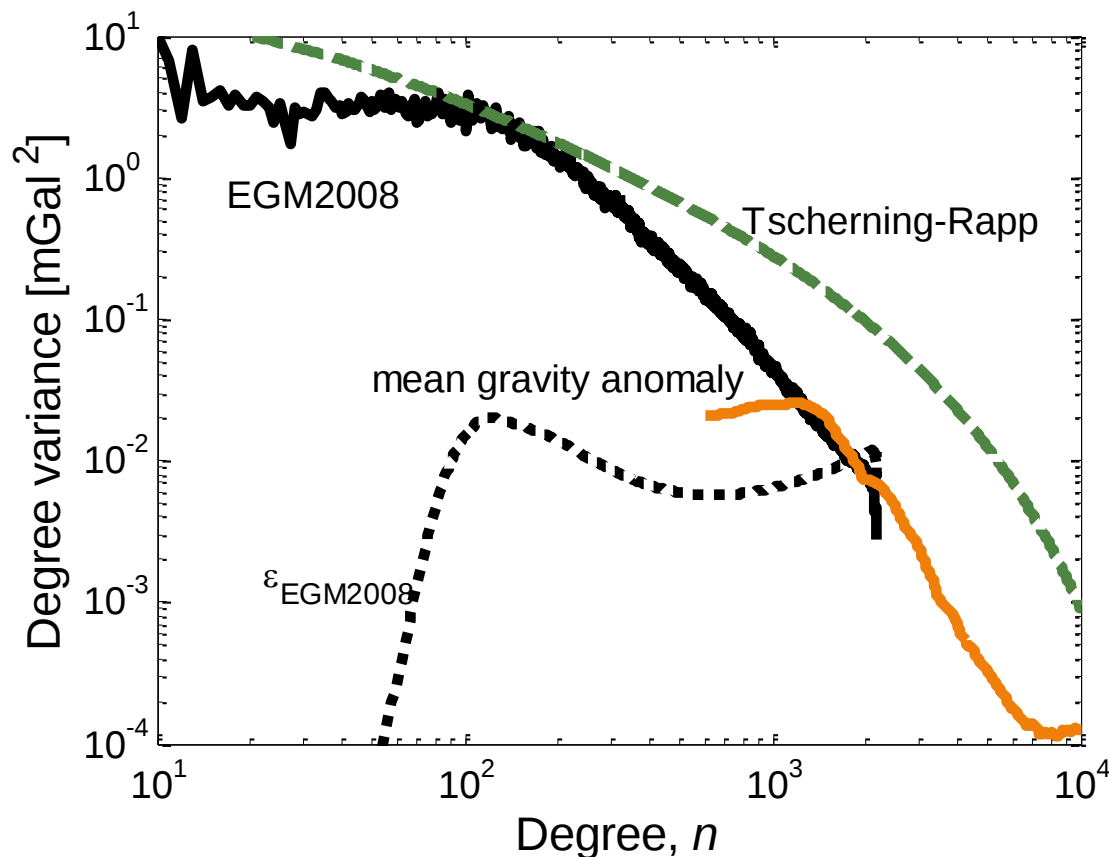
sample cut for PSD estimate
(area: 160 km x 300 km)

- Inside Hungary:
1' x 1.5' mean free-air Δg
- Outside
synthetic Δg (EGM2008 n/m
2-2160)
- RTM Δg n/m ≥ 2160
(GRAVSOFT-TC)

Statistics of RTM reduced Δg (Hungary)
[mGal]

mean	std.	min.	max.
15.0	12.8	-20.6	108.1

Gravity degree variances



- EGM2008 to d/o 600 was removed

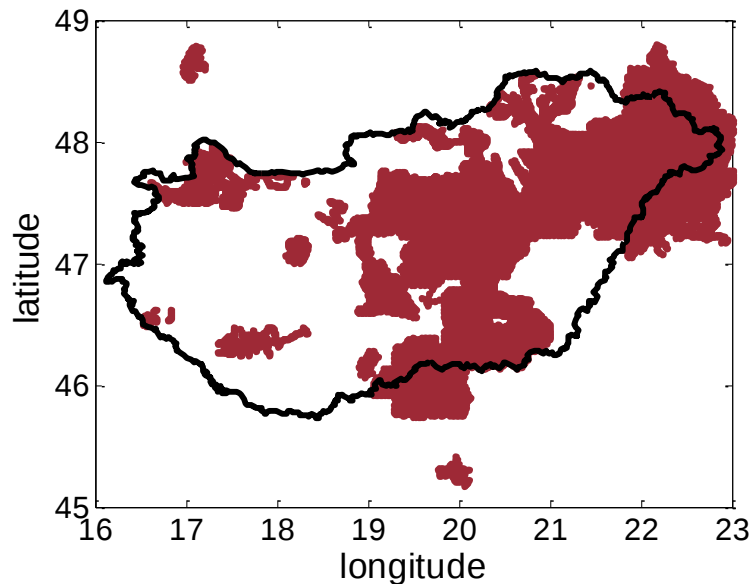
- recoverable spectral bands

$n_{\min} = 600$

$n_{\max} = 10,000$

($\sim 2\text{km}$ grid spacing)

Data sets – Horizontal gradients

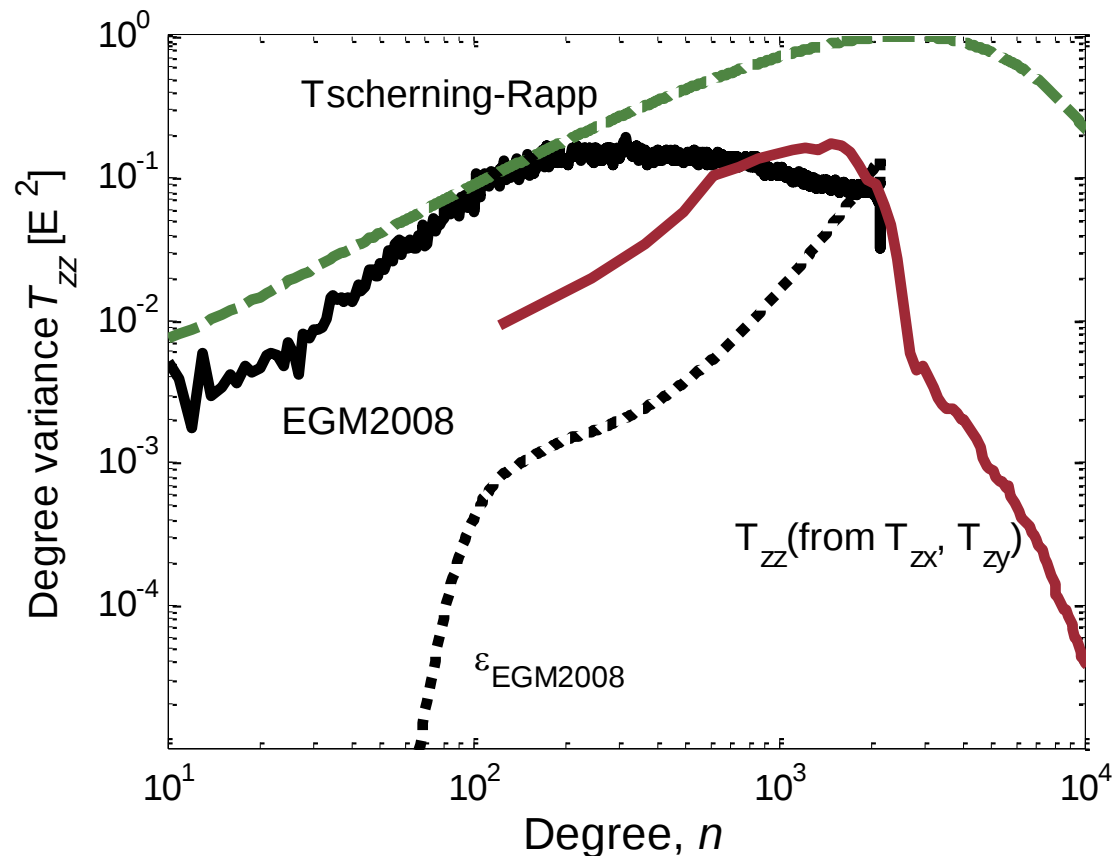


Statistics of terrain effect reduced gradients [E]

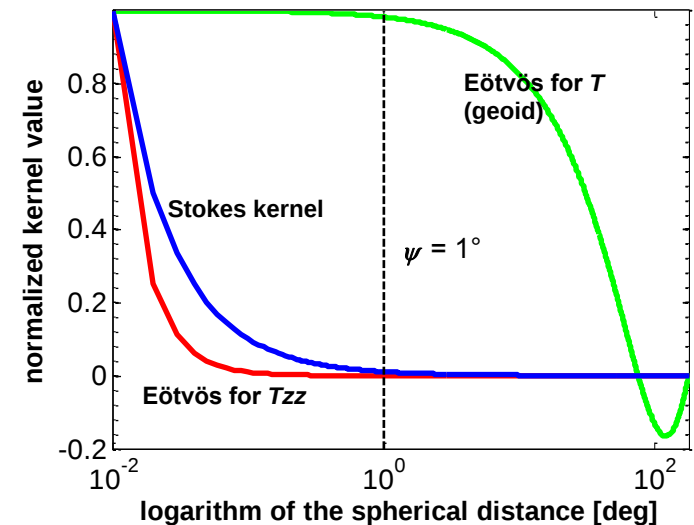
data set	mean	std.	min.	max.
T_{zy}	0.1	9.7	-88.4	72.9
T_{zx}	0.0	8.9	-67.4	71.3

- 26,000 torsion balance sites
- nearby terrain effect from levelling not RTM gradients
- grid gradients onto the grid of gravity data
- fill-in gaps: EGM2008 n/m 2- 2160

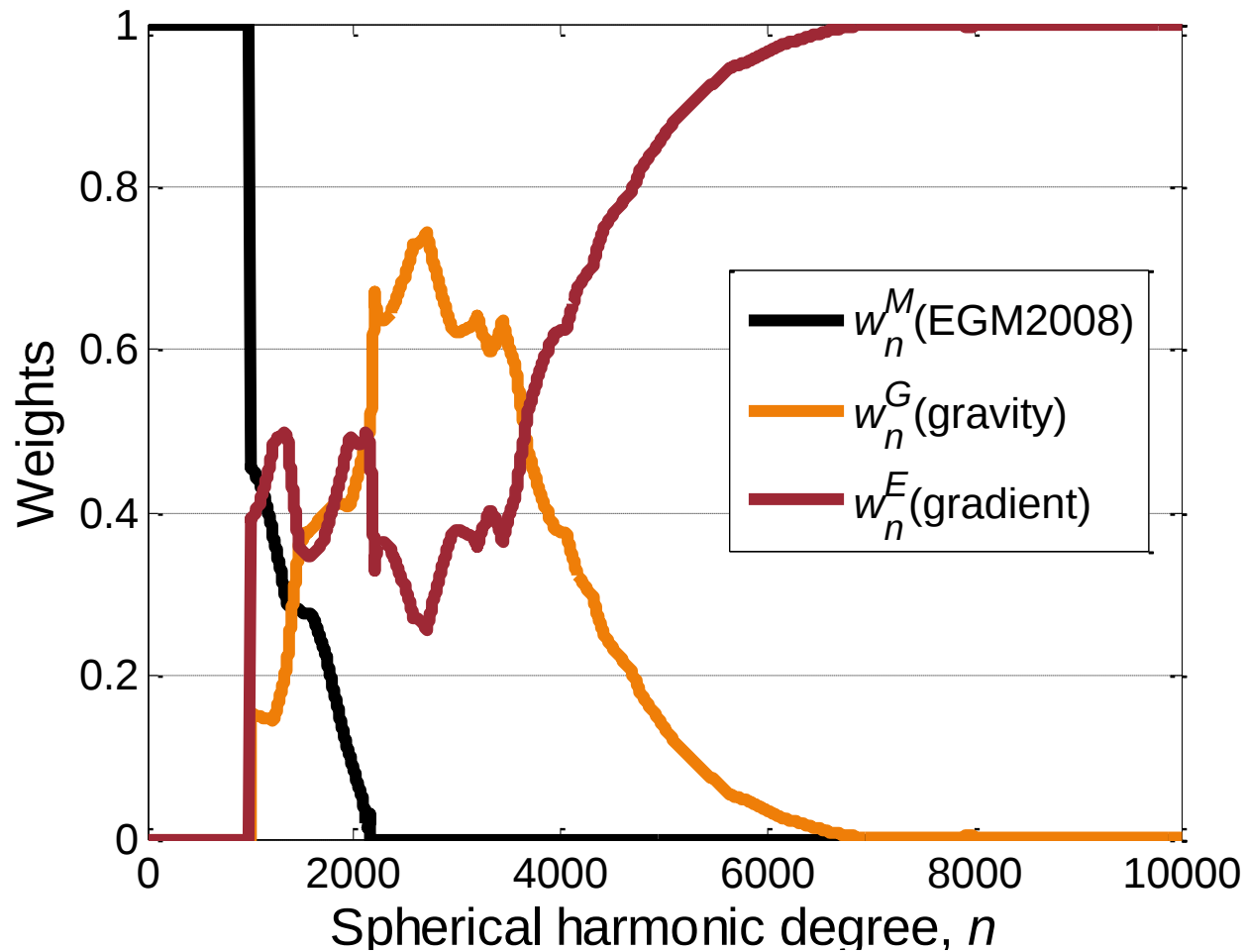
Vertical grad. degree variance



T_{zx} and T_{zy} were transformed to T_{zz} to determine spectral weights.

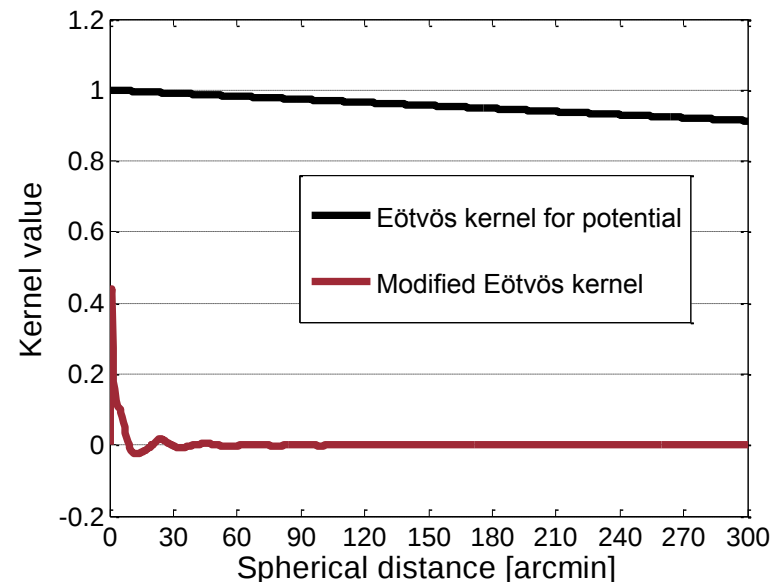
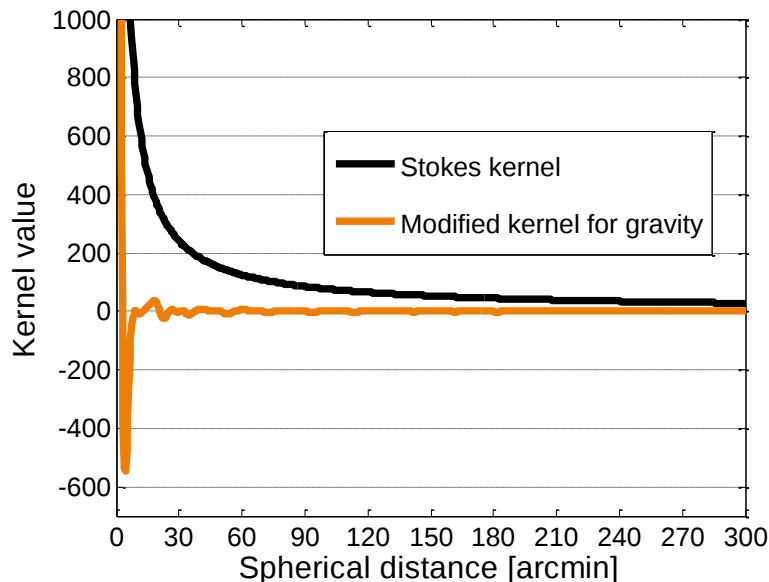


Spectral weights

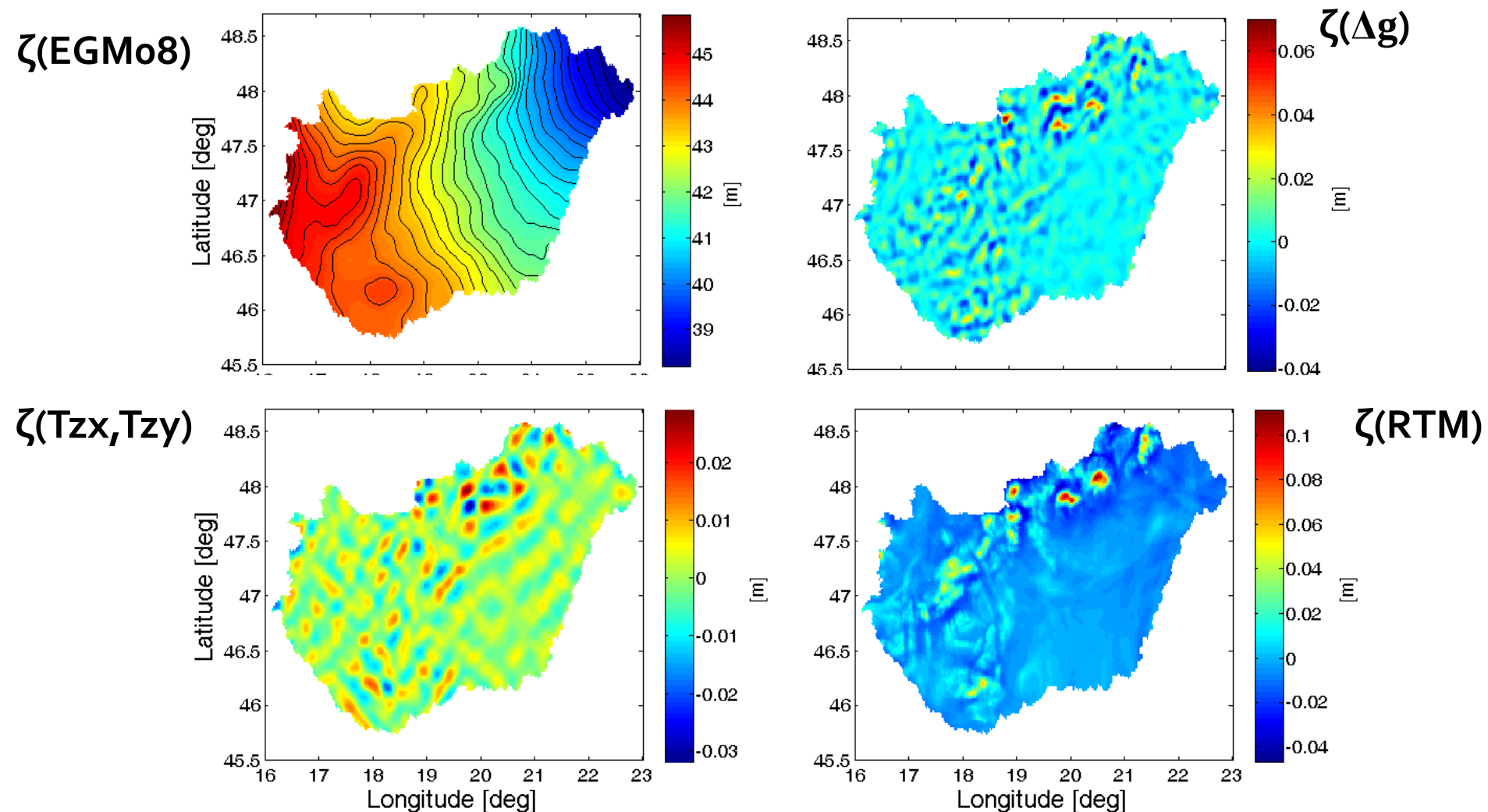


Evaluation of surface integrals

- local data solution
- 1D spherical FFT (Matlab)
- modified kernels are bounded for $\psi = 0^\circ$
- mean kernel based on Gauss-Legendre quadrature



Quasi-geoid model contributions from different data sources



Statistics of different quasi-geoid height contribution [m]

parameter	mean	std.	min.	max.
ζ (EGM2008)	42.530	1.706	38.189	45.841
ζ (gravity)	0.000	0.008	-0.040	0.067
ζ (grav. grad.)	0.000	0.005	-0.032	0.029
ζ (RTM)	-0.003	0.011	-0.047	0.111
ζ (combined)	42.527	1.709	38.175	45.868

Validation of the model

- LSC various geoid models: inconsistencies between GPS/lev. and gravimetric, astrogeodetic data sets
- re-measurement some part of the levelling network
 - significant surface subsidence (~ 10 cm, 30 years)
- Future investigations:
 - using synthetic data from EGM & topography
 - topographic effect for torsion balance data from high-resolution DTM

Mean gravity and torsion balance data sets provided by the Geological and Geophysical Institute of Hungary are gratefully acknowledged.

This study was supported by the National Scientific Research Fund
OTKA K 101603.

Thank you for your attention!