

Height and gravity changes due to glacial isostatic adjustment

Holger Steffen

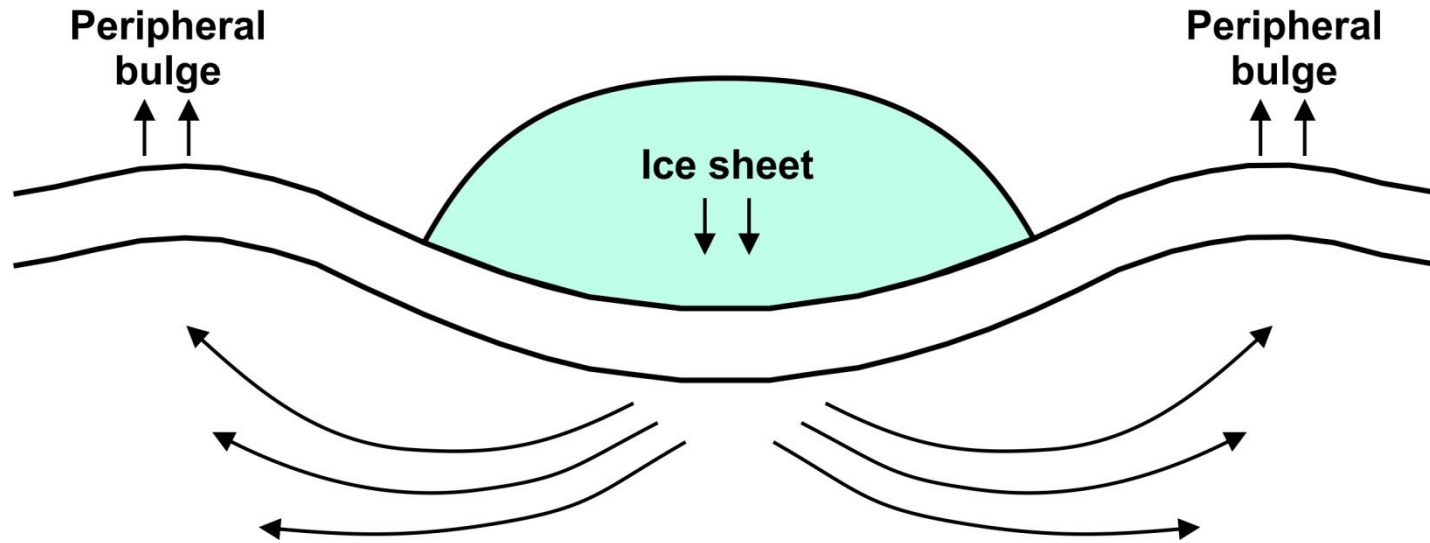
holger-soren.steffen@lm.se

L A N T M Ä T E R I E T

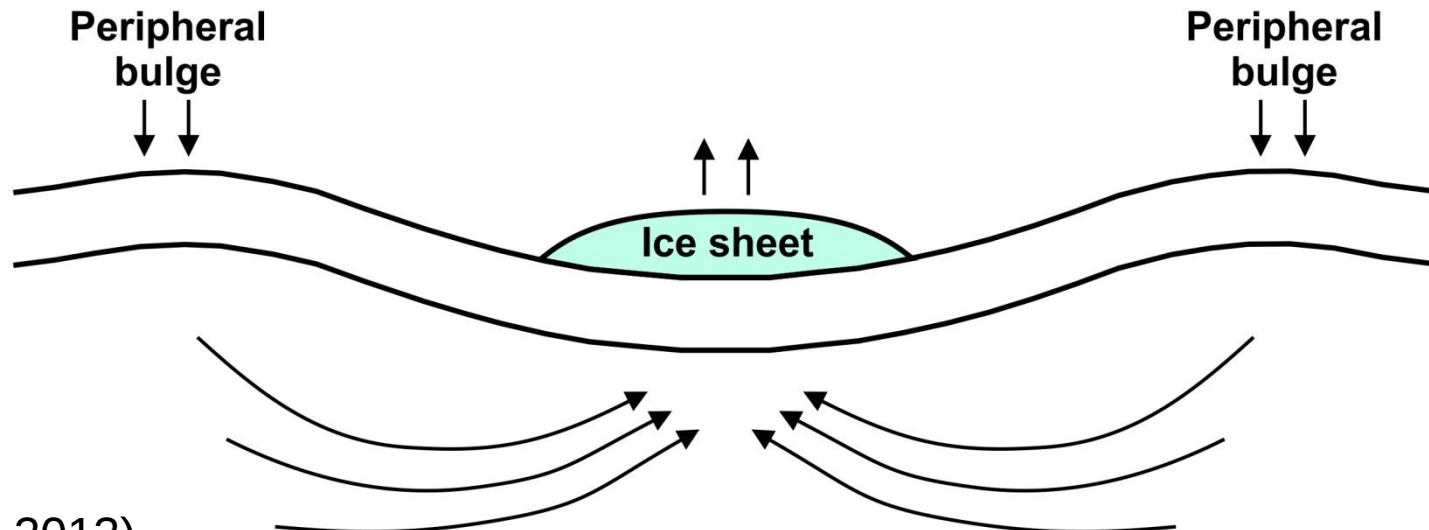


Glacial Isostatic Adjustment

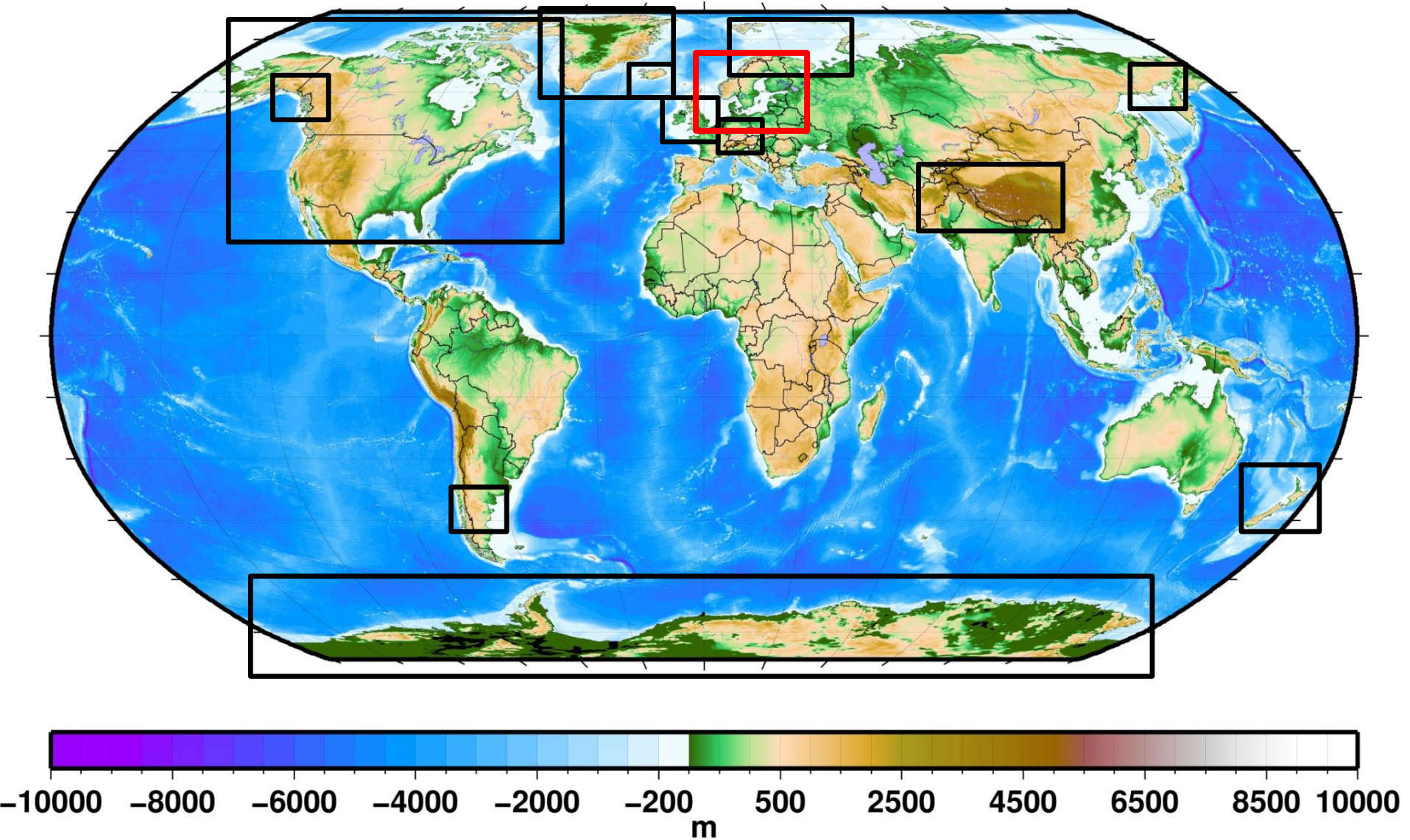
During glaciation



During deglaciation



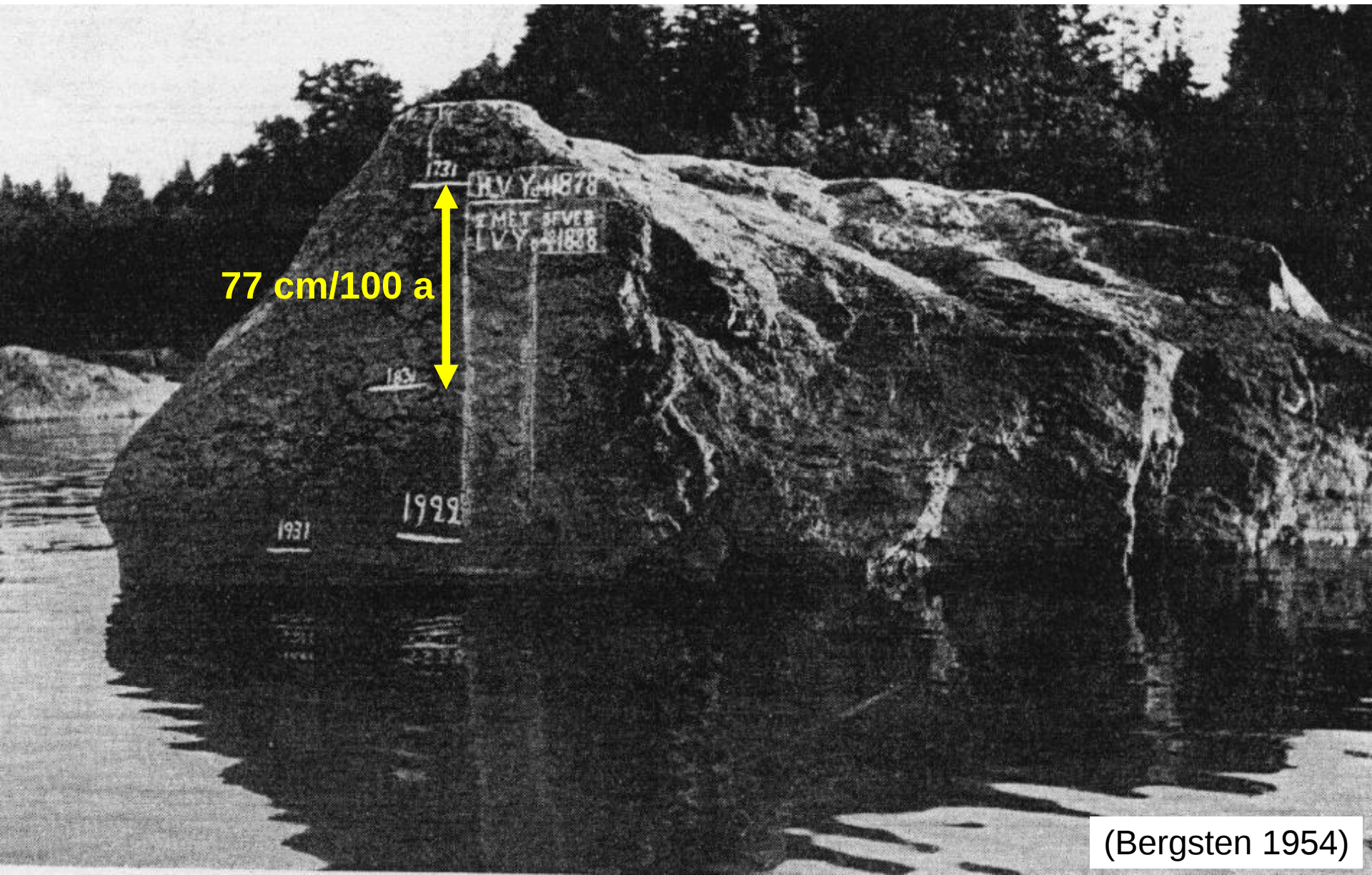
GIA around the world



GIA observations – used in modelling

- Vertical motion
 - Relative sea levels
 - Present-day rate of uplift - GPS, tide gauges, altimetry
 - Horizontal motion - GPS, VLBI, DORIS(?)
 - Gravity change due to redistribution of mass – terrestrial and space-geodetic techniques (CHAMP, GRACE)
-
- Change in Moments of Inertia
 - Polar wander
 - Non-tidal acceleration (Length Of Day)
 - Changes in the state of stress - earthquakes

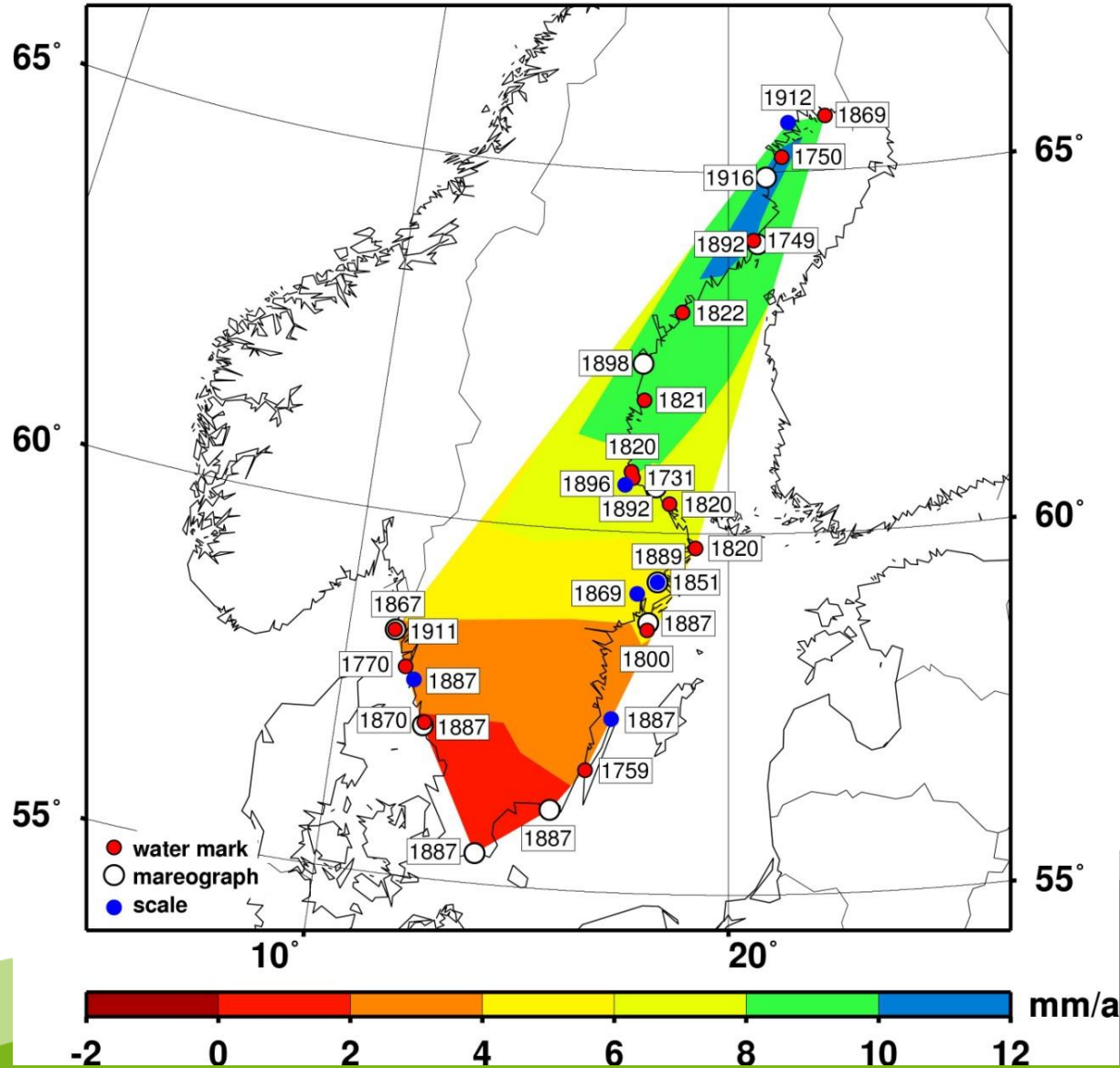
Water marks: example Celsius rock



(Bergsten 1954)

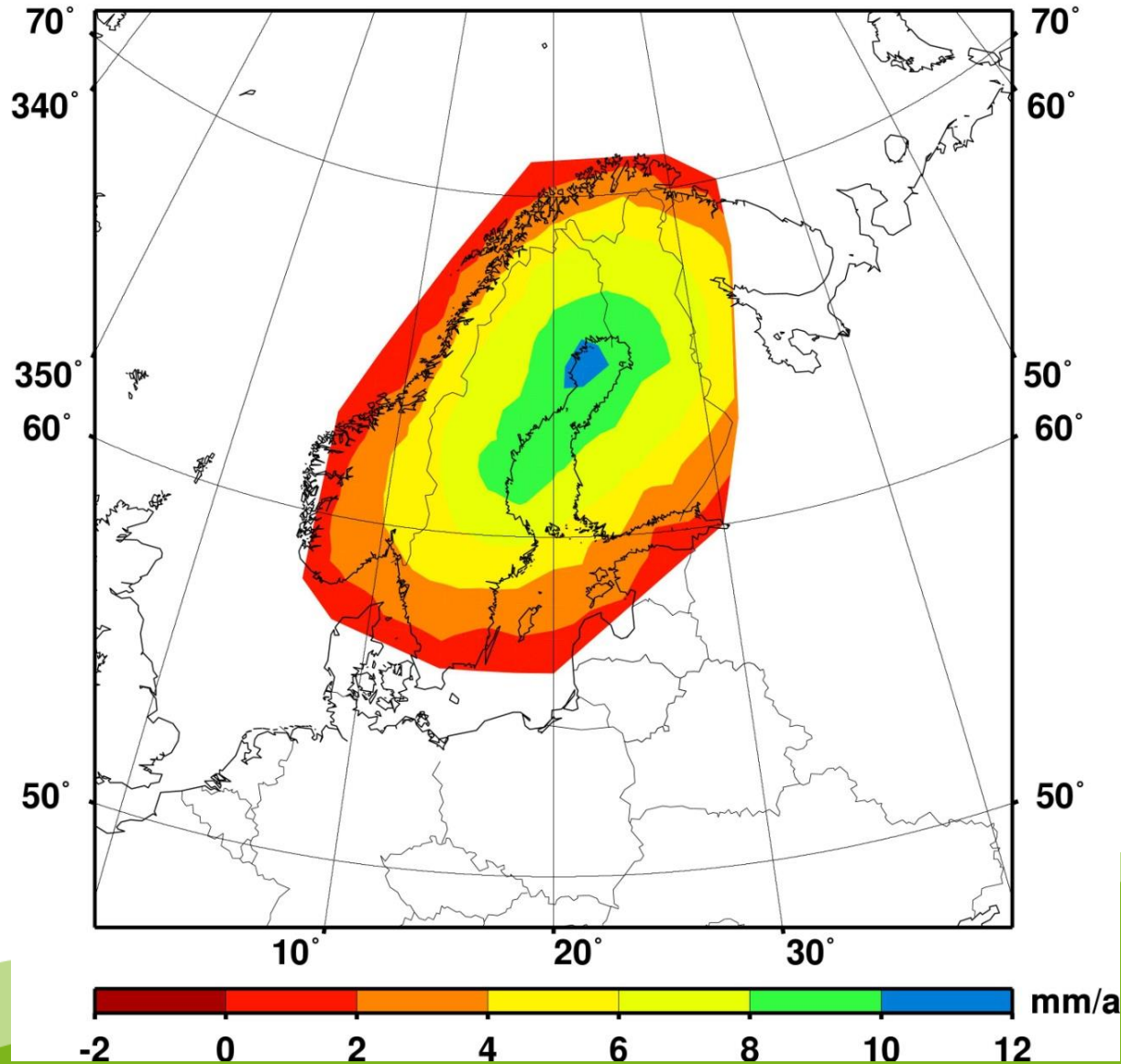
Uplift rate from water marks in Fennoscandia

Uplift from Bergsten (1954)



Uplift rate mainly from levelling

Uplift from Bjerhammar (1980)

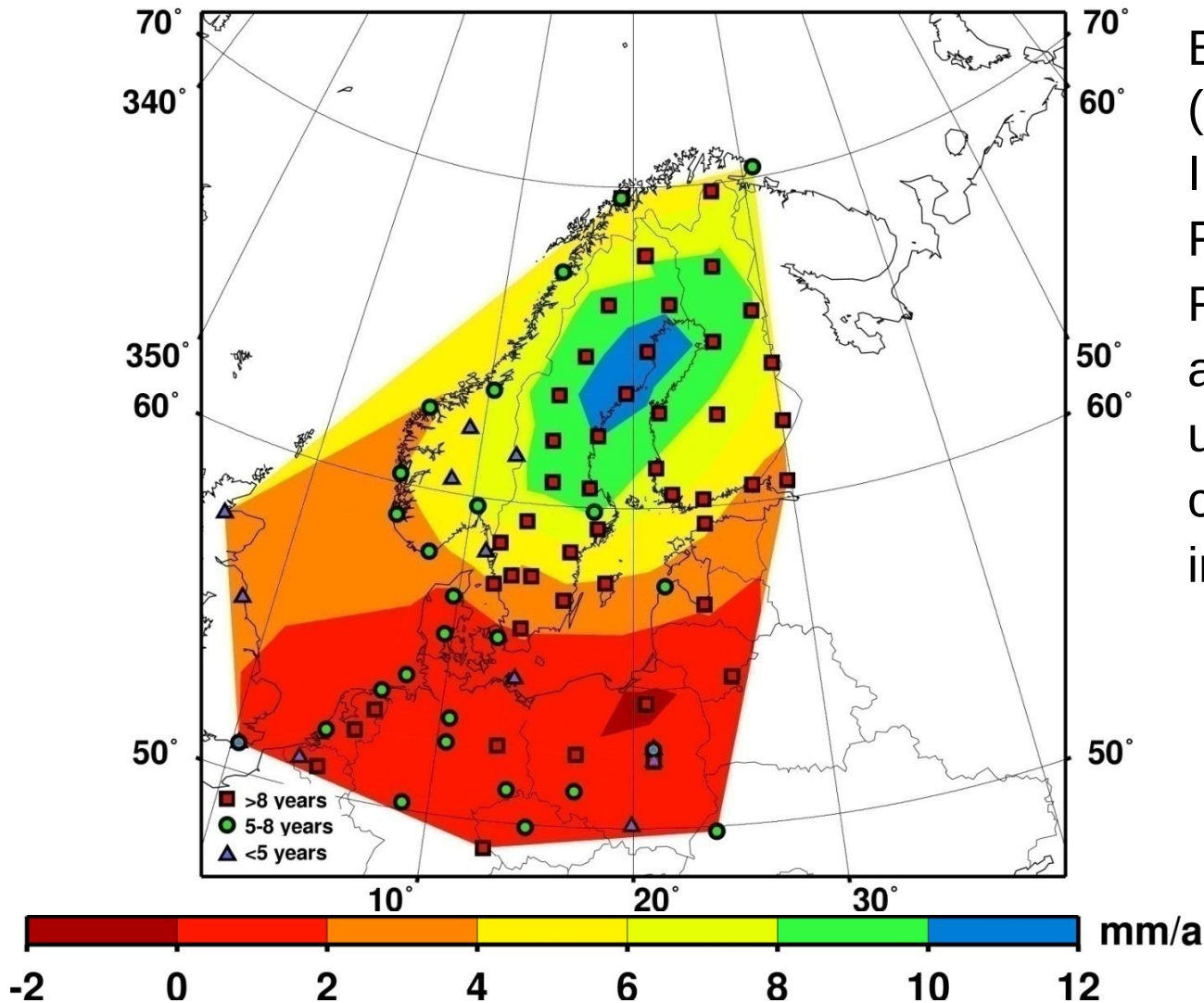


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(Steffen and Wu 2011)

GPS determined uplift rate in Fennoscandia

BIFROST (2010) observation



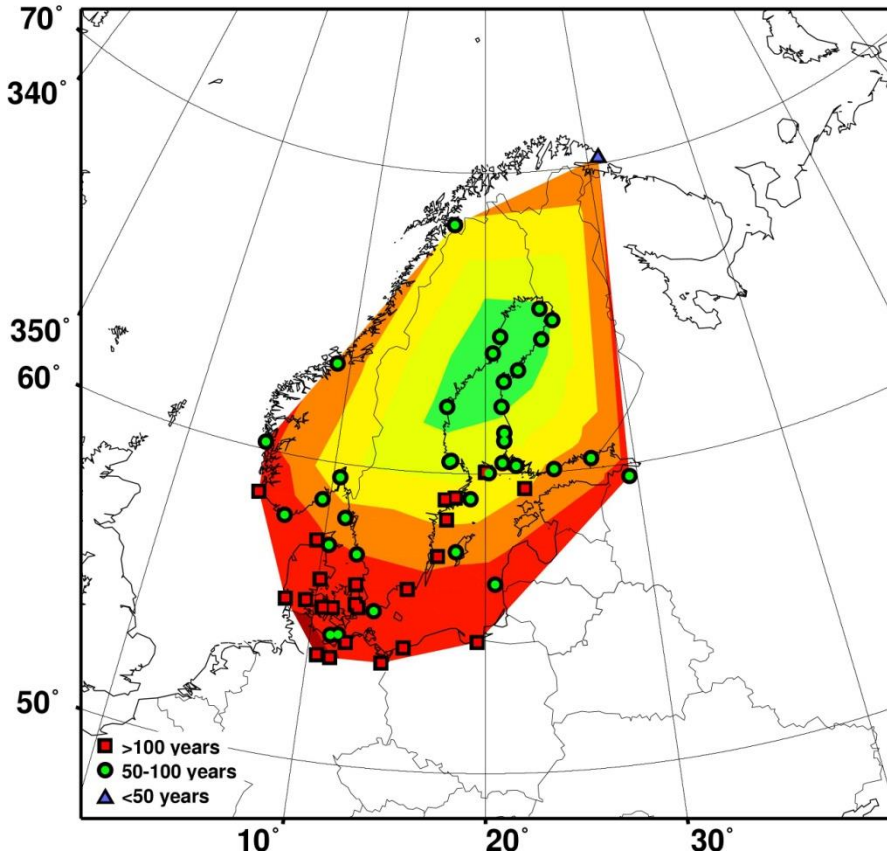
BIFROST Project
(Baseline
Inferences for
Fennoscandian
Rebound, Sea-level,
and Tectonics)
uses GPS to measure
crustal deformation
in Fennoscandia

Data from Lidberg et al. (2010)

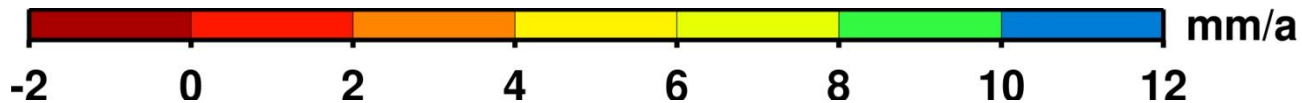
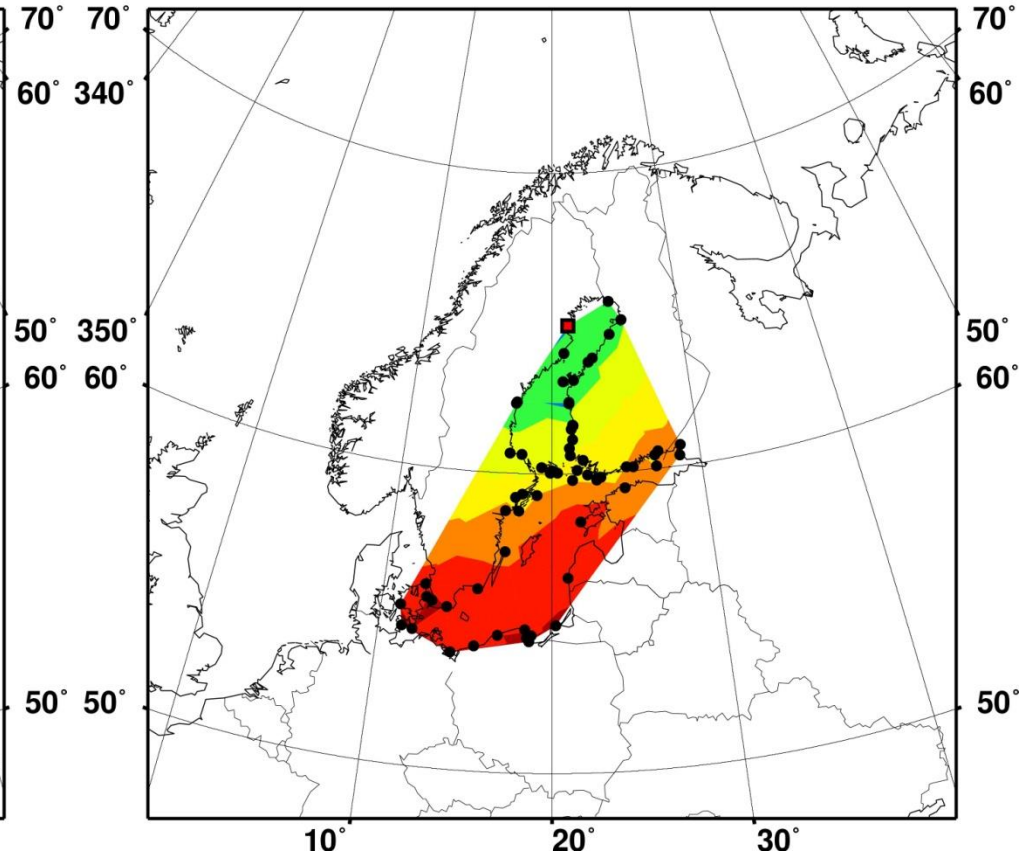
(Steffen and Wu 2011)

Tide gauge & uplift rate

Tide gauges (Ekman 1996)

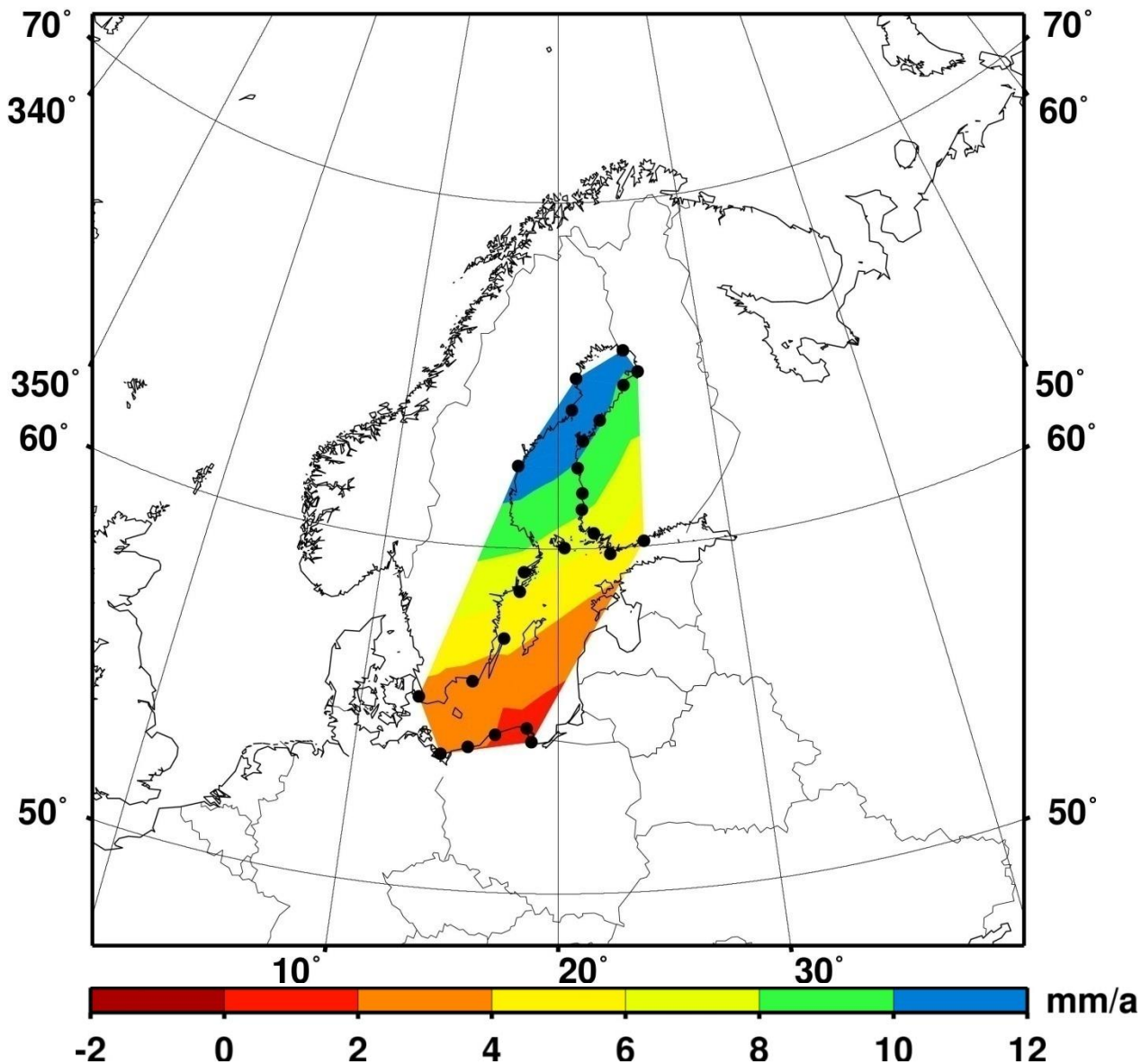


Tide gauges (Davis et al. 1999)



Satellite altimetry & uplift rate

TOPEX/POSEIDON + tide gauges (Kuo et al. 2008)



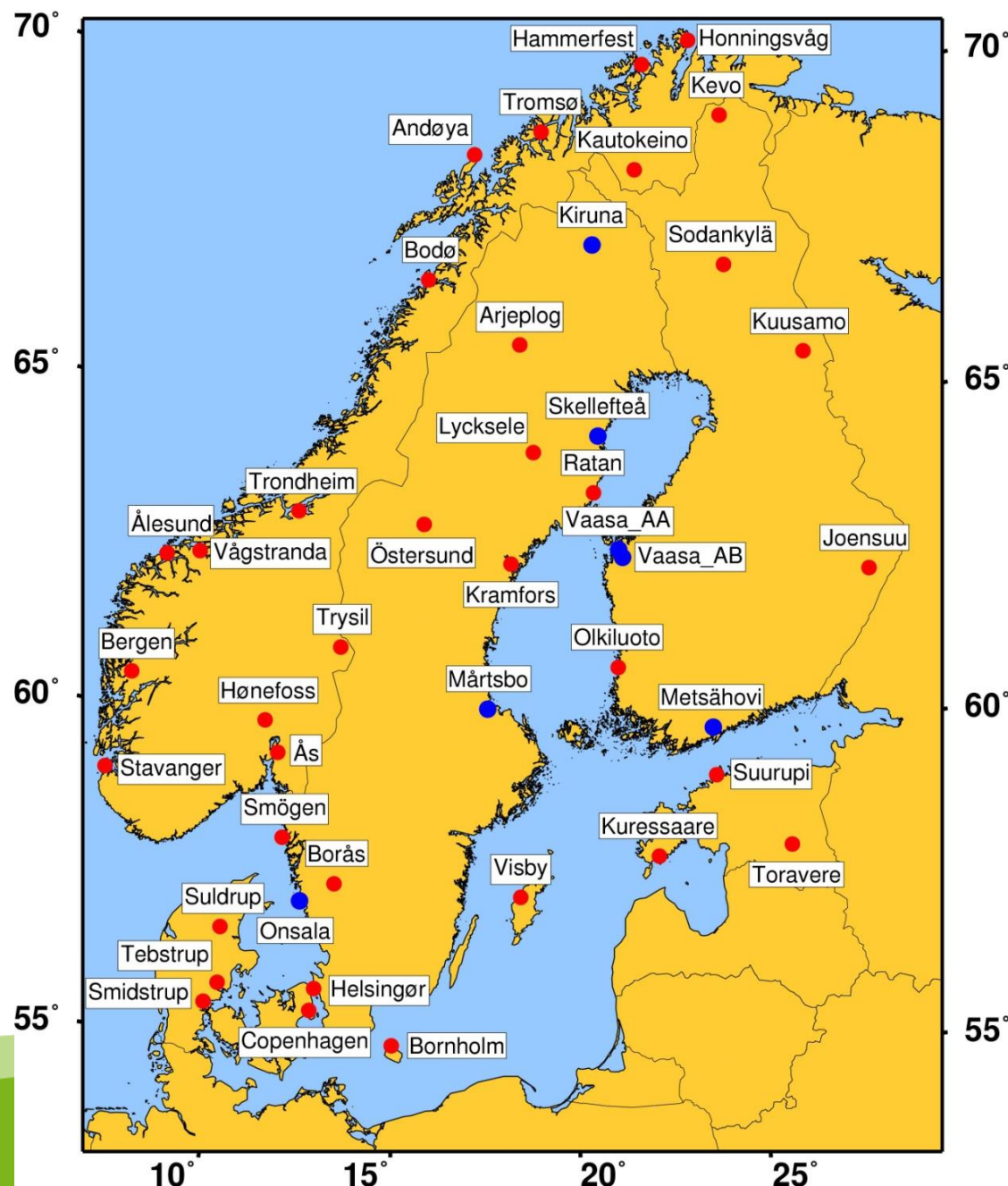
(Steffen and Wu 2011)

Absolute gravity stations in Fennoscandia



FG5-220 from IfE (Photo: Gitlein)

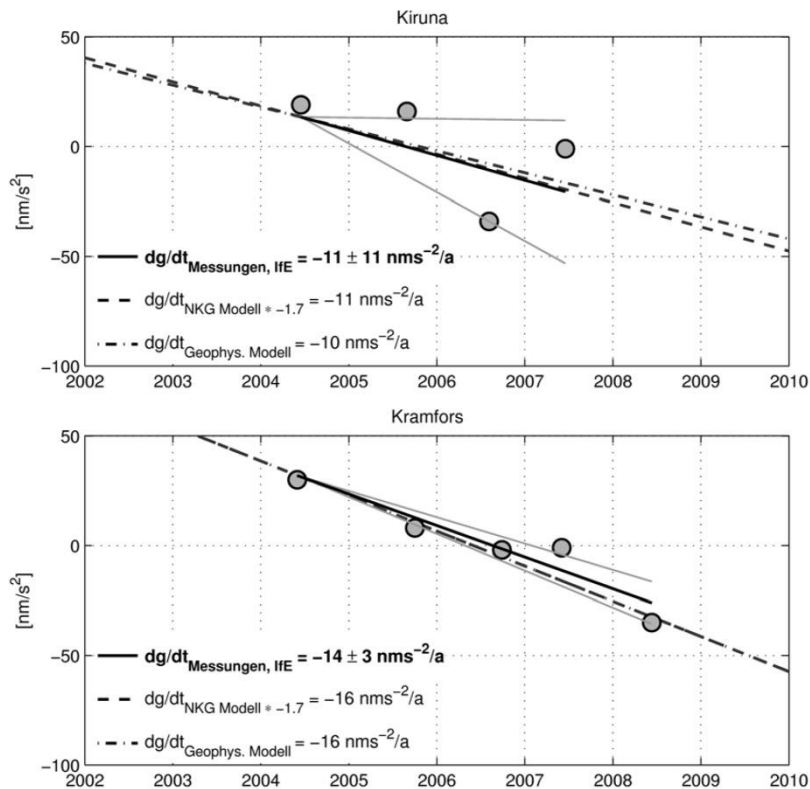
NKG stations 2003-2008



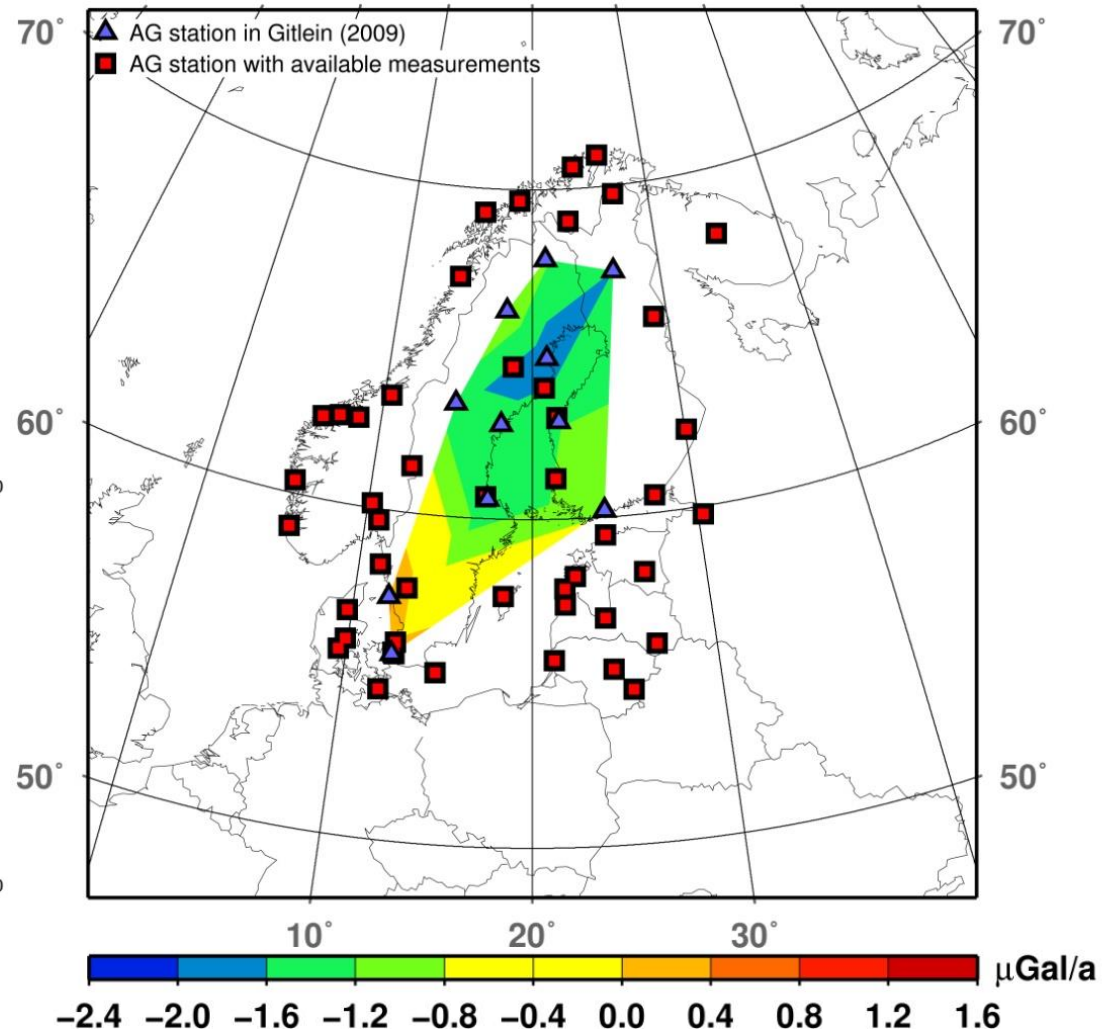
(Steffen and Wu 2011)

Vertical gravity rate from absolute gravity measurements in Fennoscandia

Gitlein (2009) result

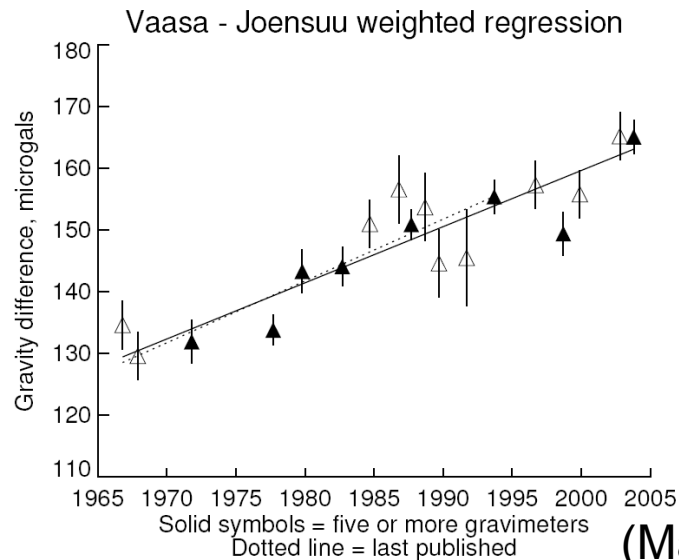
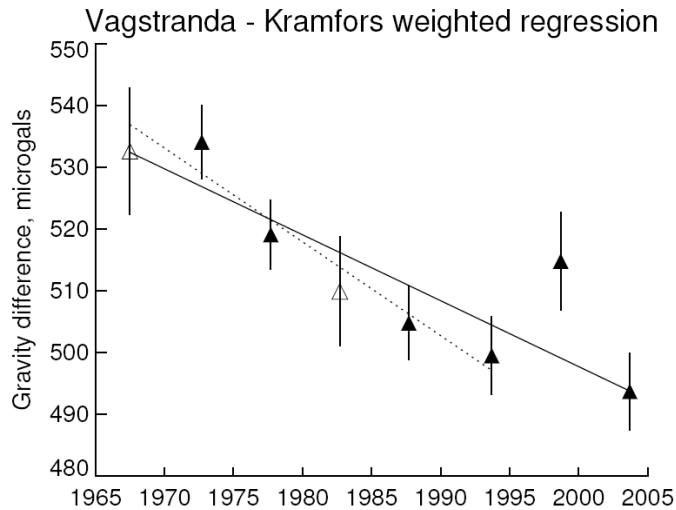


(Gitlein 2009)

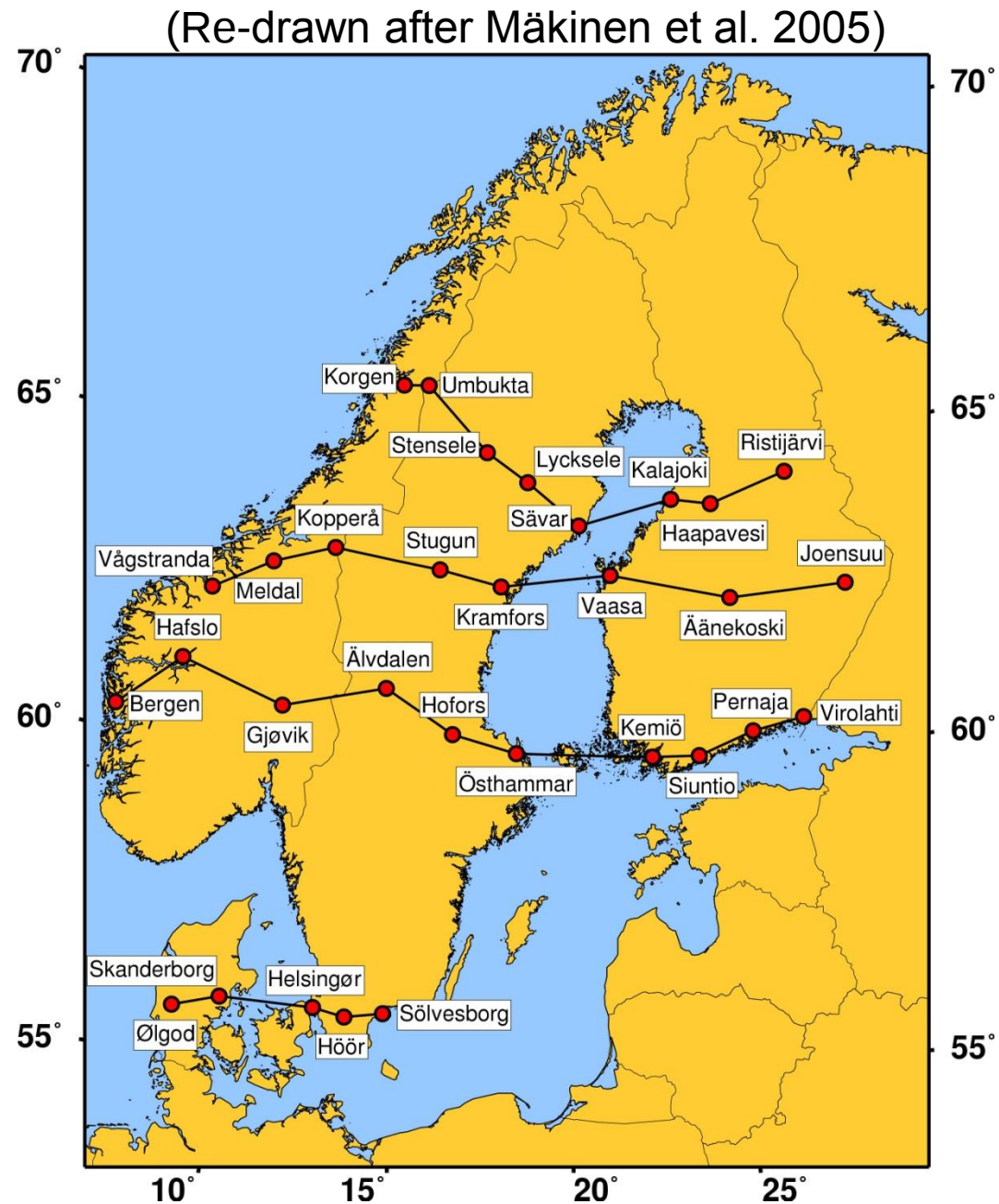


(Steffen et al. 2012, GJI)

Relative gravity measurements in Fennoscandia



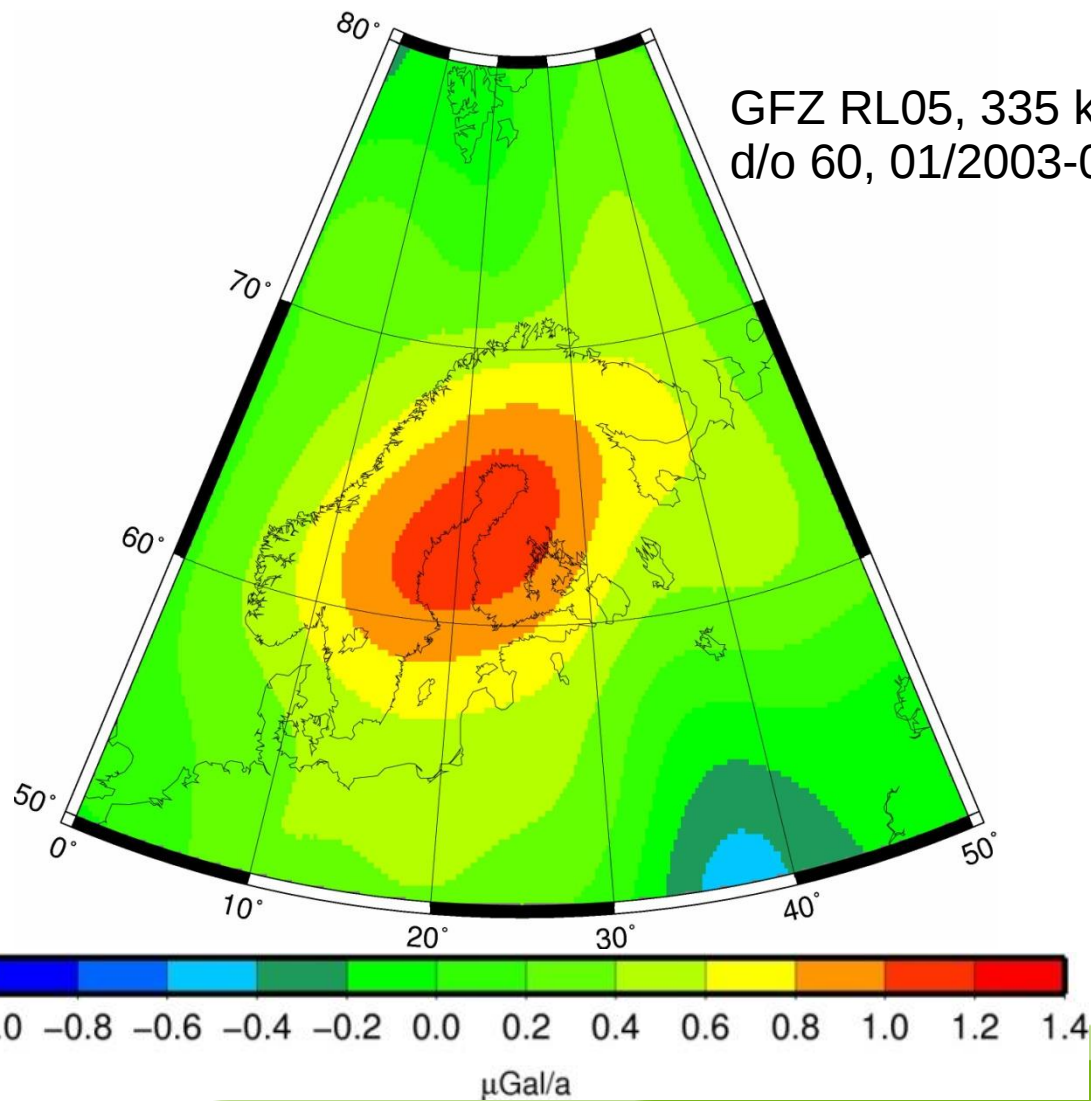
(Mäkinen et al. 2005)



(Steffen and Wu 2011)

Vertical gravity rate from GRACE

GFZ RL05, 335 km Gaussian,
d/o 60, 01/2003-02/2013



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Accuracy of GRACE for different time spans

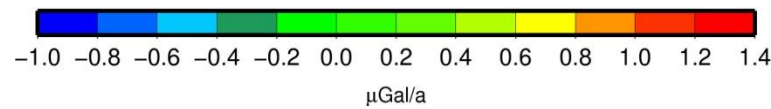
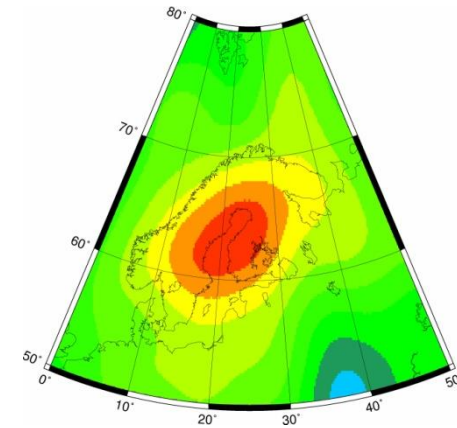
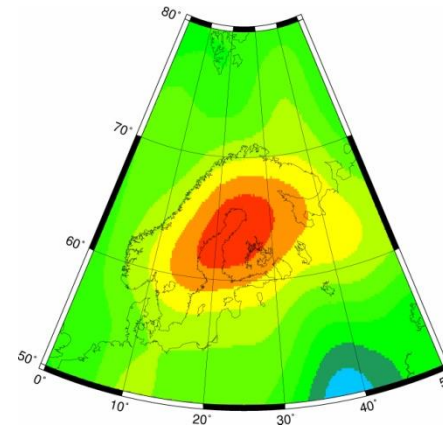
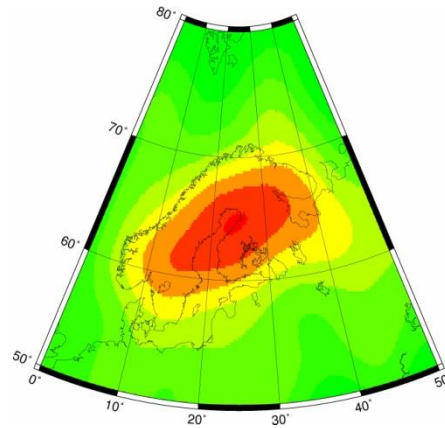
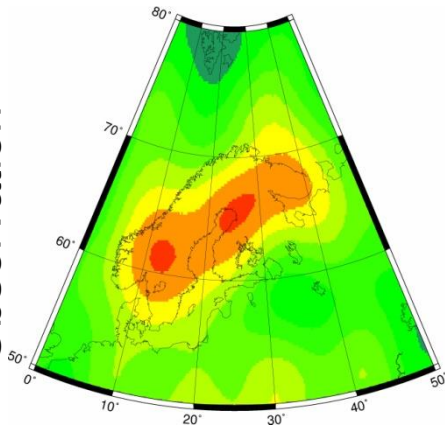
01/2003 – 12/2007

01/2003 – 12/2008

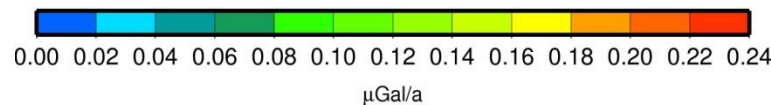
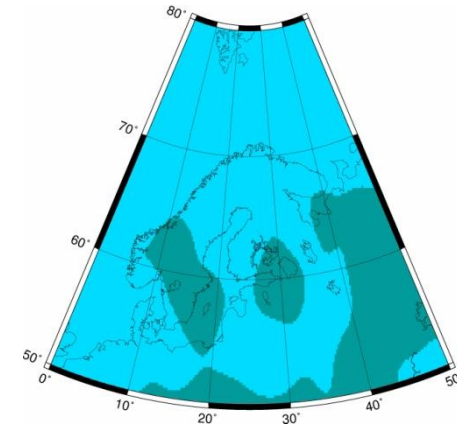
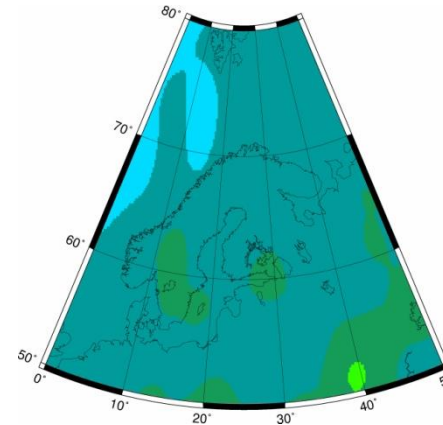
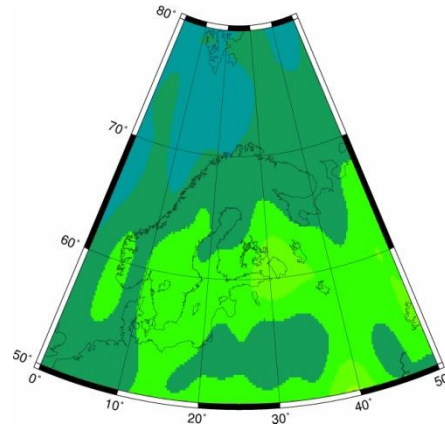
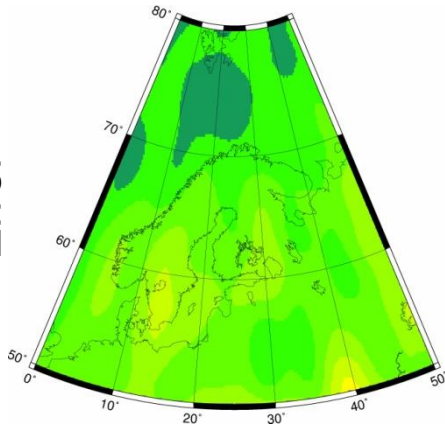
01/2003 – 12/2010

01/2003 – 12/2012

Observation



Error



Accuracy of GRACE for different time spans

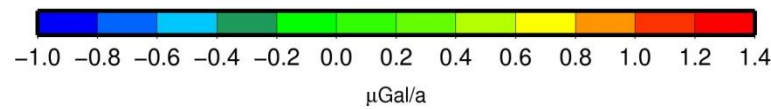
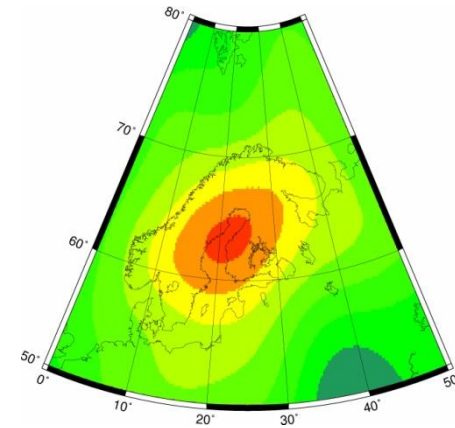
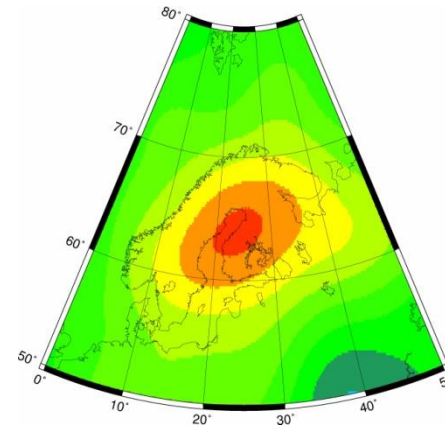
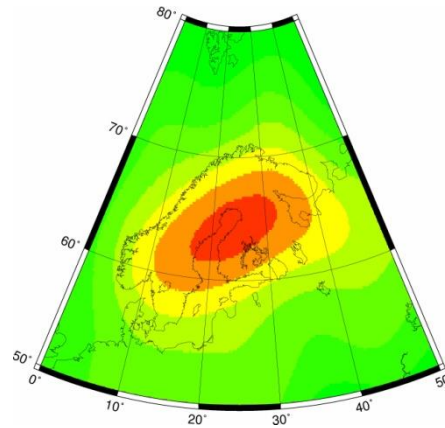
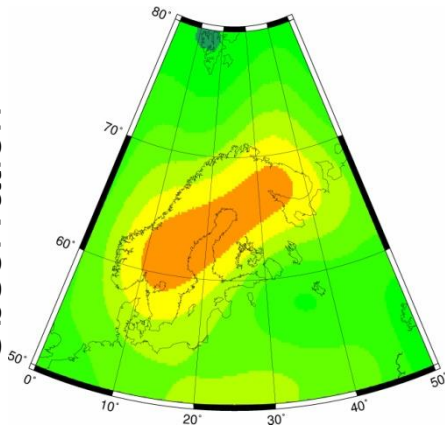
01/2003 – 12/2007

01/2003 – 12/2008

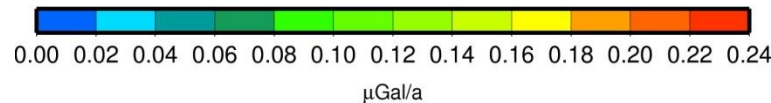
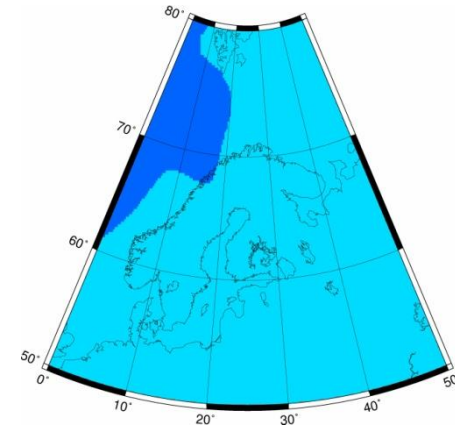
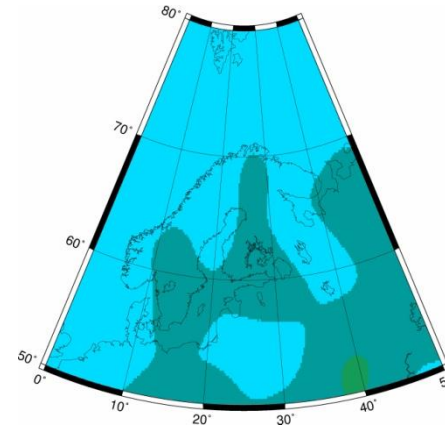
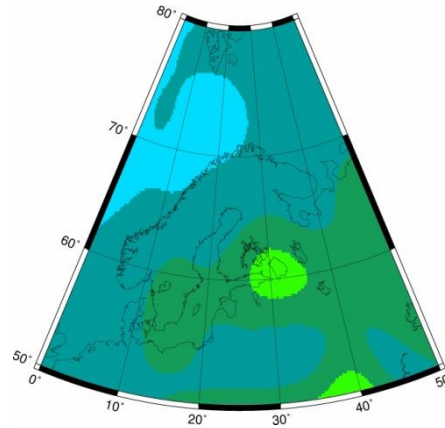
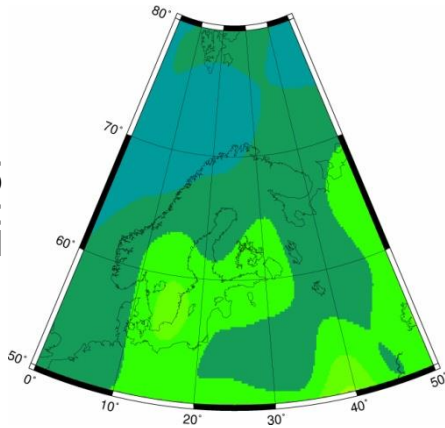
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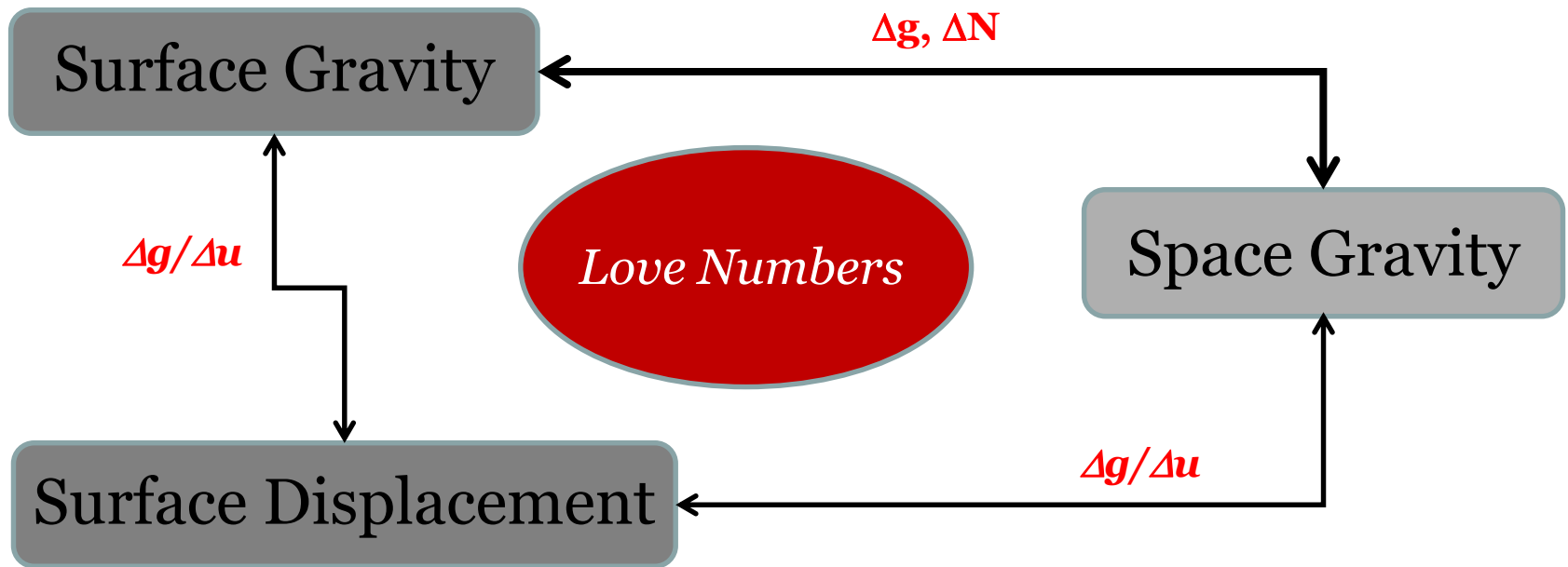
Observation



Error



Space-, ground-gravity and height changes intercomparison



Space/ground measurements

	GPS “vertical”	Absolute Gravimeter	GRACE
	Radial displacement (h_n')	Newtonian attraction (global + local (e.g. soil moisture)) + vertical displacement in gravity field (h_n') + mass redistribution (k_n')	Newtonian attraction (global) - + mass redistribution (k_n')
Resolution	mm	1-3 μGal	10 μGal 0.5 μGal
Spatial resolution	Point	Point	400 km 2,000 km
$\lambda/2$ SH coefficients	N/A	N/A	$n_{\text{max}} = 50$ $n_{\text{max}} = 10$
Temporal resolution	1 s	variable	10 days - 1 month
Long-term stability	No drift	no drift	No drift

Transfer function ratios $\Delta g/\Delta u$

$$\Delta g = \Delta g^{elastic} + \Delta g^{Newtonian} + \Delta g^{viscous}$$

(gravity variations)

$$\Delta u = \Delta u^{elastic} + \Delta u^{viscous}$$

(vertical displacement)

→ Separation of present-day ice mass changes (elastic) and (viscoelastic) glacial isostatic adjustment (post-Pleistocene deglaciation)

Viscous Transfer function $\Delta g^v / \Delta u^v$

Glacial Isostatic Adjustment (mantle density) ($n \geq 2$)

$$\frac{\Delta g^v}{\Delta u^v} = 2\pi G \rho_m \frac{n+1}{n+1/2}$$

ρ_m density of lithosphere (3.0 - 3.5 g/cm³)

(Wahr et al. 1995 GRL;
Fang & Hager 2001, GRL)
Viscoelastic Earth response

$$\approx -0.15 \mu\text{Gal/mm}$$

$$\frac{\Delta g^v}{\Delta u^v} = \frac{g_0}{a} \frac{n+1}{1.1677n - 0.5233}$$

g_0 surface gravity

a Earth's radius

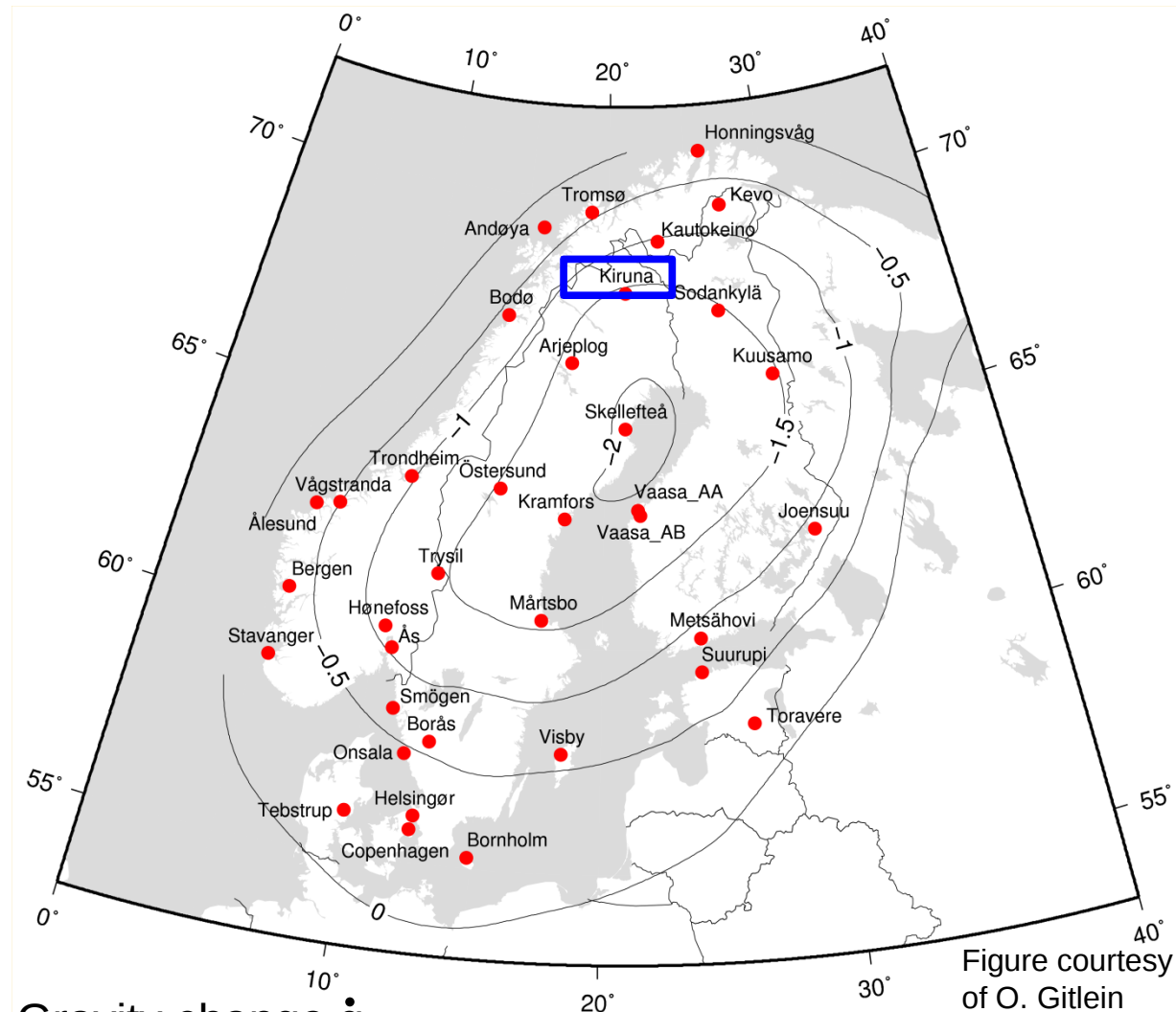
(Purcell et al. 2011, GRL)
Empirical model ($n < 60$)

**Observations: $-0.15 < \Delta g^v / \Delta u^v < -0.18 \mu\text{Gal/mm}$
(roughly constant)**

Absolute gravity network



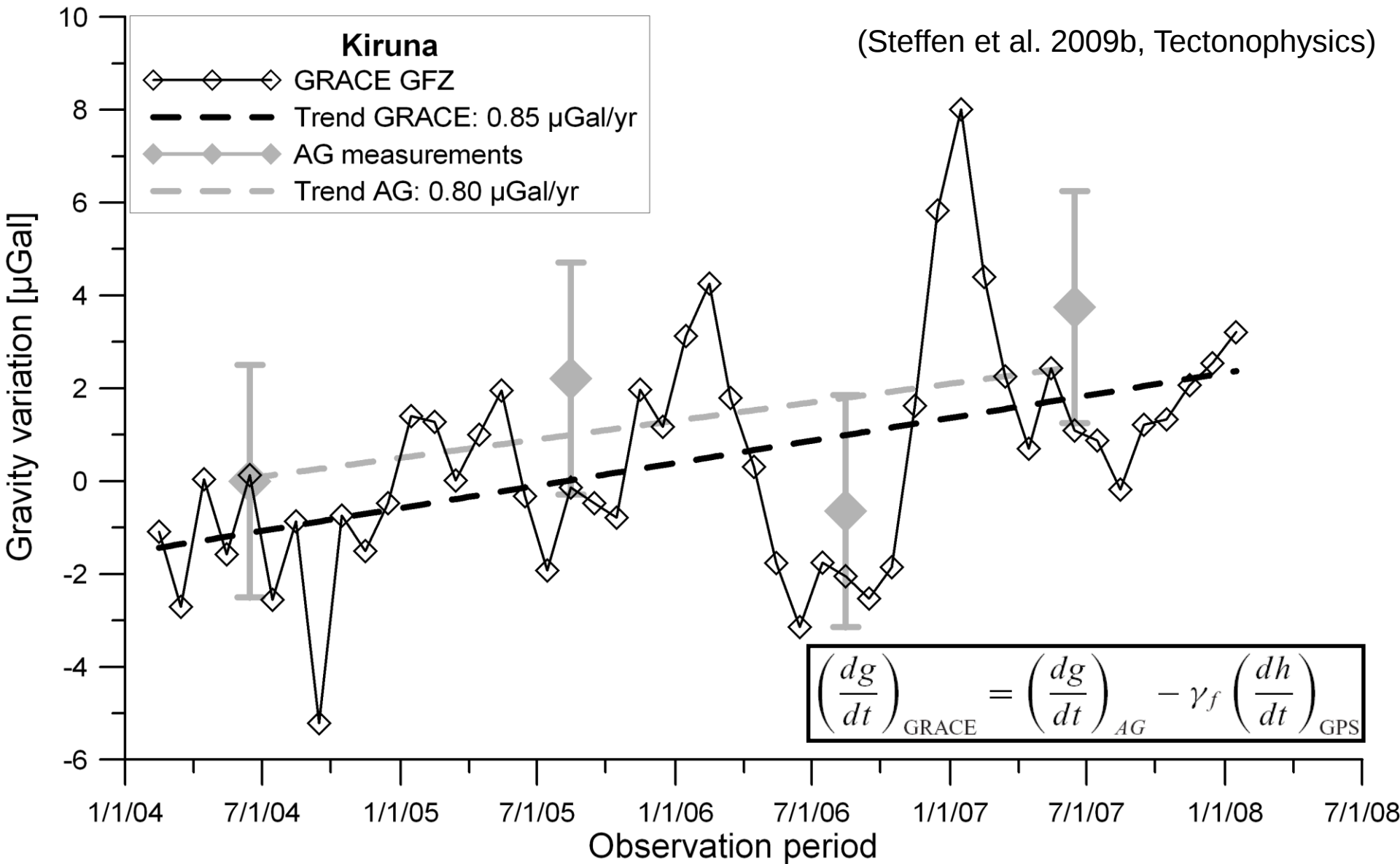
FG5-220 from IfE (Photo: Gitlein)



Gravity change $\dot{g}$ after Ekman and Mäkinen (1996)

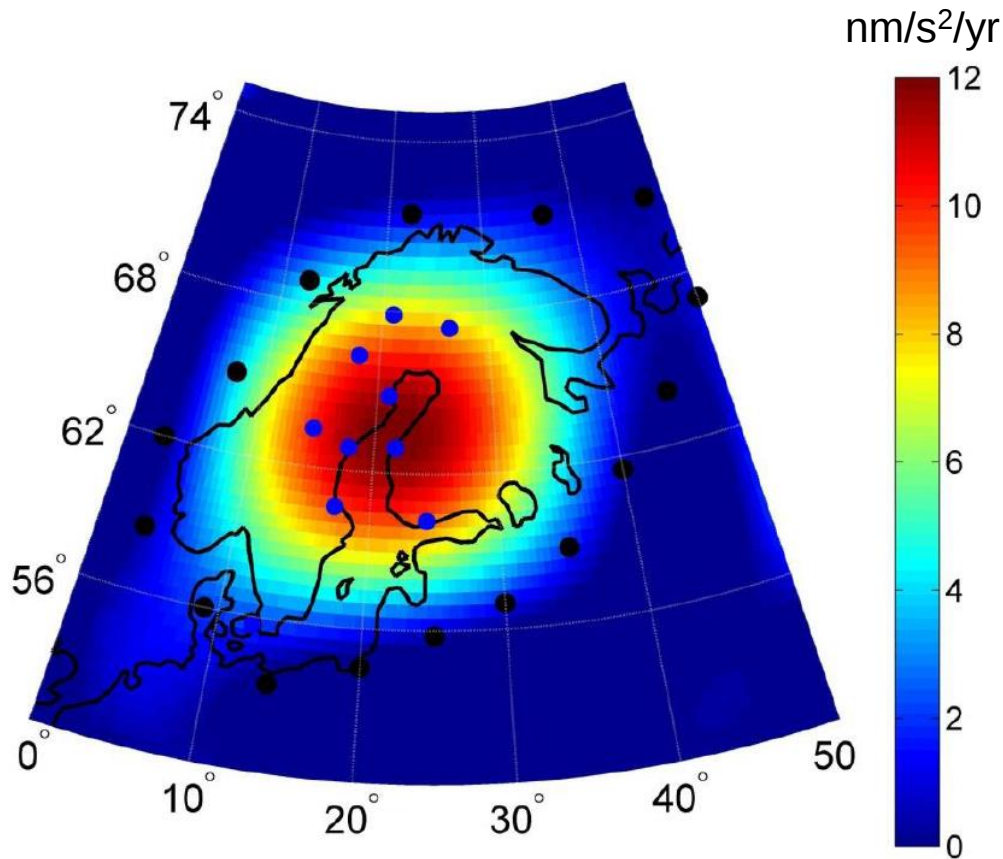


Comparison to absolute gravity

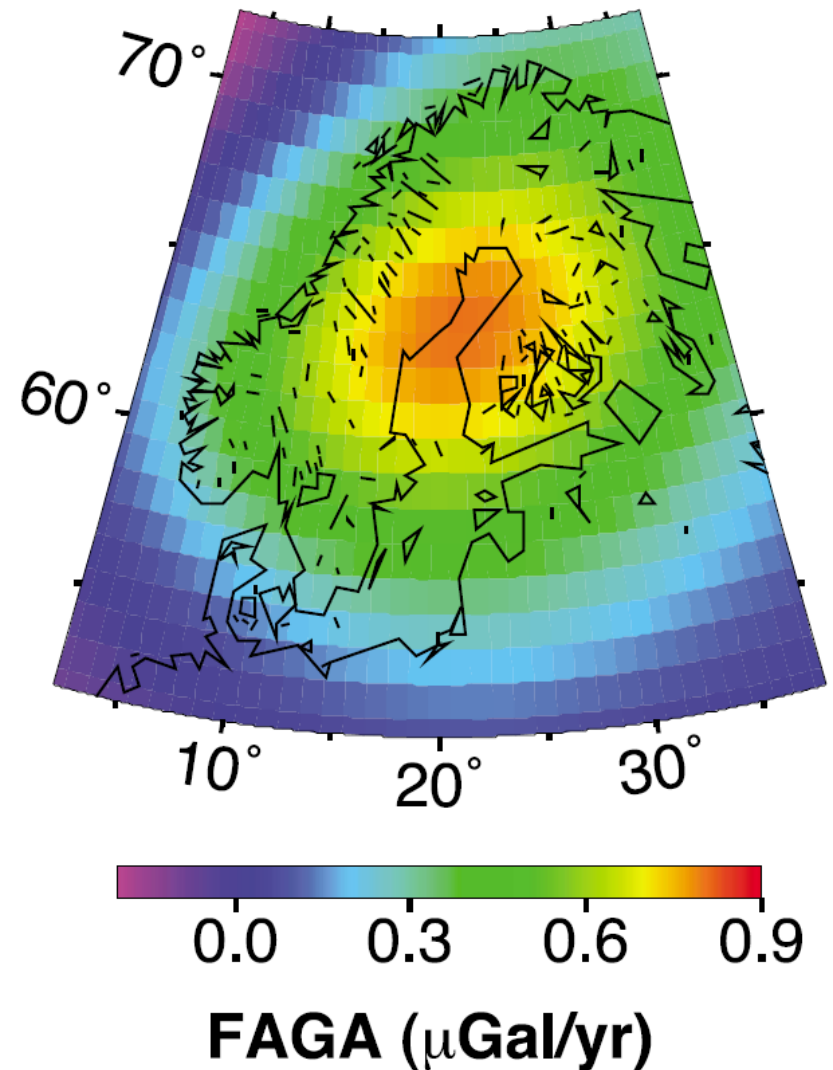


Data combination

GRACE + AG (+ GPS) +
geophysical model



GRACE + GPS + tide gauges



(Müller et al. 2012, PCE)

(Hill et al. 2011, JGR)

Conclusions

- GIA in many parts of the world
- In Fennoscandia, excellent velocity and gravity observations of high accuracy (partially long records)
 - GPS, tide gauge, levelling, absolute and relative gravimetry, GRACE
- Data needed for GIA models, which can provide inside into Earth's structure
- $\Delta g^v / \Delta h^v$ in Fennoscandia about $-0.16 - -0.17 \mu\text{Gal/mm}$
- Data combination of vertical (uplift) rates and gravity change possible – will be used to develop new GIA model for Fennoscandia

Acknowledgements

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Thank you for your attention!