Modeling surface velocities in the Southern and Eastern Alps by finite dislocations at crustal depths

A. Caporali¹ F. Neubauer² G. Stangl³ L. Ostini¹ D. Zuliani⁴

1) Department of Geosciences, University of Padova, Via Gradenigo 6, I-35131 Padova, Italy;

2)Department of Geography and Geology, University of Salzburg, Hellbrunnerstr. 34, A-5020 Salzburg,

3) AustriaAustrian Academy of Sciences and BEV, Schmiedlstr. 6, A-8042 Graz, Austria,

4) Istituto Nazionale di Oceanografia e Geofisica Sperimentale, Centro Ricerche Sismologiche, Via Treviso 55, I-31100Udine, Italy

Highlights

- Indentation of the Adria modeled by GPS data and an elastic dislocation model.
- Geometry and slip rates of six rectangular faults are constrained by least squares.
- Reverse, transpressional and normal slips from 10 to 30 mm/yr within the crust.
- Regional stress pattern from CMTs agrees with the implied strain rate eigenvectors.
- Shear heating on a half space suggests slip onset is Pliocene or more recent.

Structural setting, historical seismicity

Indentation of Adria, uplift of Tauern and lateral escape towards the Pannonian basin
Diffuse seismicity along the edge of the wedge



Available GPS velocities

Horizontal velocities of 75 permanent GPS stations resulting from the multiyear combination of the Austrian and Italian weekly solutions

BSW50 software, IGS/EPN processing standards

Reduction of the velocities to a common Eurasian frame by subtraction of rigid rotation about a common pole



Elastic dislocation model

- Measured velocities are the effect of slip at depth along 6 suitable planes
- Use geology to set up the planes
- Change the plane parameters until the rss of the observed velocities minus modelec velocities is minimal

Fault id.	Name	Long.	Lat.	depth	strike	dip	right lat.	reverse	Length	Width	Momentrate
		(deg)	(deg)	(km)	(deg)	(deg)	(m/yr)	(m/yr)	(km)	(km)	10 ¹⁸ J/yr
1	Giudicarie	11.04	46.31	30	211	80	-0.01	0.03	58	10	0.56
2	NAWC	12.18	47.85	10	84	60	-0.01	0.02	297	12	2.02
3	Pustertal	12.91	46.68	20	282	89	0.01	0.01	123	30	1.41
4	TW north	12.22	47.34	8	261	89	0.00	0.02	61	20	0.55
5	Brenner fault	11.49	47.02	2	188	45	0.00	-0.02	35	4	0.08
6	<u>Dinarids</u>	14.13	45.62	5	312	45	0.01	0.01	122	14	0.73





The surface kinematics predicted (white arrows) by the best fitting model



Predicted vertical motion (mm/yr)

We show the vertical motion predicted by the elastic model
The actual vertical motion includes also effects of GIA and erosion, which are not modeled in this work

Hence comparison with vertical uplift data should be made with care.
Blue polygons represent Individual seismic sources from DISS



T_{0,} q₀



Frictional heating at the fault plane generates considerable heat flow, A model for the surface temperature and heat flow can be set up, depending on a number of parameters The surface temperature and heat flow can be measured and used to constrain the

heating process at depth

Conclusion: for shear stress in the range 100 – 300 Mpa, assuming no lateral fluid advection and nominal diffusion coefficients of the rocks, the absence of high heat flow at the surface implies that the slip process must be relatively recent (Plio-Pleistocene) relative to the time of collision of Adria with the Southern Alps (Oligocene-Miocene)

Conclusion

- GPS velocities can be modeled in deformation areas to constrain geophysical processes at depth
- With the geophysical model we can predict the velocities at any point at the surface of the deformed area
- The Eastern/Southern Alps could then be an example in support to the WG on Geokinematics (chair: M. Lidberg)