



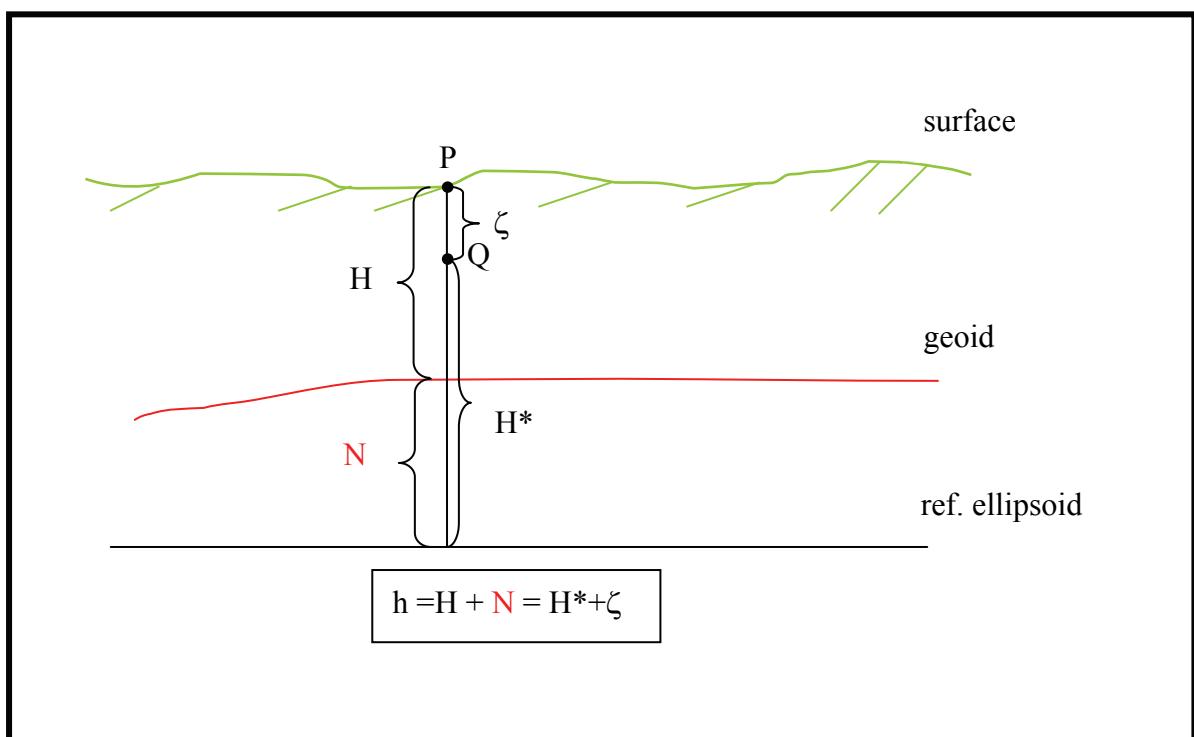
# A Rigorous Formula For Geoid-to-Quasigeoid Separation

by

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# 1 Introduction

The traditional formula:

$$\mathbf{N} - \zeta = \mathbf{H}^* - \mathbf{H} = \frac{\bar{g} - \bar{\gamma}}{\bar{\gamma}} \mathbf{H}, \quad (1)$$

where  $\bar{g}$  and  $\bar{\gamma}$  are the mean gravity between the geoid and Earth's surface and mean normal gravity between the reference ellipsoid and normal height. Also

$\mathbf{N} - \zeta$  = geoid-quasigeoid heights,

and

$\mathbf{H}^* - \mathbf{H}$  = normal height - orthometric height.

**Problem:** Convert (1) to a more practical form.

The standard practical formula:

$$\mathbf{N} - \zeta \approx \frac{\Delta g_p^B}{\bar{\gamma}} \mathbf{H}, \quad (2)$$

where  $\Delta g_p^B$  = (simple) Bouguer gravity anomaly.

**How can we improve Eq. (2)?**

## 2 Basic assumptions

### Assumptions and notations:

#### 1. Bruns' formula

$$\zeta = \frac{T}{\gamma} \quad (3)$$

holds for any point  $P'$  between the geoid and the Earth's surface (including these surfaces),

**T = disturbing potential**

**$\gamma$  = normal gravity at normal height vs.  $P'$ .**

### Examples:

$$N = \frac{T_g}{\gamma_0} = \text{geoid height} \quad (4a)$$

and

$$\zeta_P = \frac{T_p}{\gamma_Q} = \text{quasigeoid height} \quad (4b)$$

where  $T_g$  and  $T_p$  are the disturbing potentials at the geoid and surface point P, respectively, and  $\gamma_0$  and  $\gamma_Q$  are the normal gravity values at the reference ellipsoid and normal height, respectively

## 2. The no-topography and topographical quasigeoid heights are defined by

$$\zeta^{nt} = \frac{\mathbf{T} - \mathbf{V}^t}{\gamma} = \frac{\mathbf{T}_g^{nt}}{\gamma} \quad \text{and} \quad \zeta^t = \frac{\mathbf{V}^t}{\gamma}, \quad (5)$$

where  $\gamma$  is located at normal height vs. the position of  $\mathbf{T}$  and  $\mathbf{V}^t$ .  $\mathbf{T}^{nt}$  is the no-topography disturbing potential.

## 3. The quasigeoid height, the disturbing potential and the gravity anomaly (all denoted with superscript $i=0$ ) can be decomposed into the no-topography ( $i = nt$ ) and topographic ( $i=t$ ) components, e.g.

$$\zeta = \zeta^0 = \zeta^{nt} + \zeta^t. \quad (6)$$

## 4. The gravity anomalies are related to the disturbing potentials according the approximation provided by the “boundary condition”(cf. Heiskanen and Moritz 1967, p. 85 )

$$\Delta \mathbf{g}^j = -\frac{\partial \mathbf{T}^j}{\partial h} + \frac{\mathbf{T}^j}{\gamma} \frac{\partial \gamma}{\partial h}, \quad j=0, nt, t. \quad (7)$$

where  $h$  is the height along the normal to the reference ellipsoid.

**Note 1.** The no-topography gravity anomaly differs from the refined Bouguer gravity anomaly as follows:

$$\Delta \mathbf{g}^{nt} = \Delta \mathbf{g}^{\text{BO}} - \frac{\mathbf{V}^t}{\gamma} \frac{\partial \gamma}{\partial \mathbf{h}} \quad (8a)$$

where

$$\Delta \mathbf{g}^{\text{BO}} = \Delta \mathbf{g} - \mathbf{A}^t, \quad \text{with} \quad \mathbf{A}^t = -\frac{\partial \mathbf{V}^t}{\partial \mathbf{h}}. \quad (8b)$$

**Note 2.** Eqs. (3) and (7) imply that

$$\left( \frac{\partial \zeta^j}{\partial \mathbf{h}} \right) = \frac{\partial}{\partial \mathbf{h}} \left( \frac{\mathbf{T}^j}{\gamma} \right) = - \left( \frac{\Delta \mathbf{g}^j}{\gamma} \right); \quad j = 0, nt, t. \quad (9)$$

### 3 The main results

**Proposition 1:**  $\mathbf{N} - \zeta_P = \int_0^{H_p} \frac{\Delta \mathbf{g}}{\gamma} d\mathbf{h}. \quad (10)$

**Proof:** The proposition follows directly from Eqs. (9):

$$\mathbf{N} \cdot \boldsymbol{\zeta}_P = \int_{H_P}^0 \frac{\partial \boldsymbol{\zeta}}{\partial \mathbf{h}} d\mathbf{h} = \int_0^{H_P} \frac{\Delta \mathbf{g}}{\gamma} d\mathbf{h}. \quad (11)$$

**Proposition 2:**

$$\mathbf{N} \cdot \boldsymbol{\zeta}_P = \frac{\Delta \mathbf{g}_P^{BO}}{\bar{\gamma}} \mathbf{H}_P + \mathbf{T}\mathbf{C} + \mathbf{G}\mathbf{C}, \quad (12a)$$

where

$$\mathbf{T}\mathbf{C} = \underbrace{\frac{\mathbf{V}_g^t}{\gamma_0} - \frac{\mathbf{V}_P^t}{\gamma_Q}}_{TC1} - \frac{\mathbf{V}_P^t \mathbf{H}_P}{\gamma_Q \bar{\gamma}} \frac{\partial \gamma}{\partial \mathbf{h}}, \quad (12b)$$

(topographic correction)

and

$$\mathbf{G}\mathbf{C} = \int_0^{H_P} \left( \frac{\Delta \mathbf{g}^{nt}}{\gamma} - \frac{\Delta \mathbf{g}_P^{nt}}{\bar{\gamma}} \right) d\mathbf{h}. \quad (12c)$$

(gravimetric correction)

**Proof:** The proof is given below.

## Proof of Proposition 2

**By introducing the notation**

$$TC1 = \frac{V_g^t}{\gamma_0} - \frac{V_p^t}{\gamma_Q}, \quad (P1)$$

**and noting that**

$$\frac{\Delta g^{nt}}{\bar{\gamma}} = \frac{\Delta g^{BO}}{\bar{\gamma}} - \frac{V_p^t H_p}{\gamma_Q \bar{\gamma}} \frac{\partial \gamma}{\partial h}, \quad (P2)$$

**it follows from Prop. 1 that the right member of the proposition can be written**

$$RM = TC1 + \int_0^{H_p} \frac{\Delta g^{nt}}{\gamma} dh = TC1 + \int_{H_p}^0 \frac{\partial}{\partial h} (\zeta^{nt}) dh = TC1 + N^{nt} - \zeta_P^{nt} = N - \zeta_P, \quad (P3)$$

**and the proposition follows.**

## 4 The topographic and gravimetric corrections

### 4.1 The topographic correction

- The topographic correction can be written

$$TC = T\bar{C} + dTC , \quad (13a)$$

where

$$\bar{TC} = \frac{V_g^t - V_P^t}{\bar{\gamma}} \quad (13b)$$

and

$$dTC = V_g^t \left( \frac{1}{\gamma_0} - \frac{1}{\bar{\gamma}} \right) - V_P^t \left( \frac{1}{\gamma_Q} - \frac{1}{\bar{\gamma}} \right) - \frac{V_P^t H_P}{\gamma_Q \bar{\gamma}} \frac{\partial \gamma}{\partial h} \approx \frac{V_g^t - V_P^t}{2\bar{\gamma}^2} H_P \frac{\partial \gamma}{\partial h}. \quad (13c)$$

Considering the spherical approximation (cf. Heiskanen and Moritz 1967, p. 87)

$$\frac{1}{\bar{\gamma}} \frac{\partial \gamma}{\partial h} \approx -\frac{2}{R}, \quad (14)$$

where R is sea-level radius, Eq. (13c) becomes

$$dTC \approx -\frac{V_g^t - V_P^t}{\bar{\gamma}} \frac{H_P}{R}. \quad (15)$$

As  $H_p/R < 1.4 \times 10^{-3}$ ,  $\mathcal{D} T b$  can be safely neglected in practical applications of Eq. (13a), and the topographic correction thus agrees with that of Flury and Rummel (2009.)

- Decomposing the topographic potential into the potentials of a Bouguer plate and a terrain effect ( $dV^t$ ), and noting that the Bouguer plate potential is the same on top as at the bottom of the plate, Eq. (13b) becomes

$$\bar{T}\bar{C} = \frac{dV_g^t - dV_P^t}{\bar{\gamma}}, \quad (16)$$

i.e. the topographic effect reduces to the difference between the terrain effects at the geoid and the surface point.

## 4.2 The gravimetric correction

The strict gravimetric correction, Eq. (11c), can be rewritten and developed into a Taylor series as follows:

$$GC = - \int_0^{H_p} \left[ \frac{\partial \zeta^{nt}}{\partial h} + \frac{\Delta g_P^{nt}}{\bar{\gamma}} \right] dh = \zeta_g^{nt} - \zeta_P^{nt} - \frac{\Delta g_P^{nt}}{\bar{\gamma}} H_p = \\ \sum_{k=2}^{\infty} \frac{H_p^k}{k! \gamma_Q} (-1)^k \left( \frac{\partial \zeta^{nt}}{\partial h} \right)_P^k + \Delta g_P^{nt} H_p \left( \frac{1}{\gamma_Q} - \frac{1}{\bar{\gamma}} \right) \approx - \frac{H_p^2}{2 \gamma_Q} \left( \frac{\partial \Delta g^{nt}}{\partial h} \right)_P, \quad (17)$$

where we have truncated the series to its first term

## 5 Conclusions

- The strict formula was given by Proposition 2.
- A practical formula reads

$$N - \zeta \approx \frac{\Delta g_P^{BO}}{\bar{\gamma}} H_P + \frac{dV_g^t - dV_P^t}{\bar{\gamma}} - \frac{H_P^2}{2\gamma_Q} \left( \frac{\partial \Delta g^{nt}}{\partial h} \right)_P, \quad (18)$$

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**THANK YOU!**