The Abruzzo earthquake: temporal and spatial analysis of the first geodetic results

L. Biagi, S. Caldera, D. Dominici, F. Sansò

Politecnico di Milano Università degli studi de L'Aquila

Outline

Introduction: the Abruzzo earthquake

The network and the processing strategies

Time modelling of daily results: displacements estimation at earthquake epoch

Spatial interpretation of the horizontal displacements

The vertical displacements

Future works



Abruzzo earthquake

Main event: 6th april, 1:33 UTC Location: 42.33° N, 13.33° E, Depth: 8.8 km Magnitude: 5.8 Richter



Before and after the main event: many other pre seismic and after seismic events

Abruzzo earthquake

Main event: 6th april, 1:33 UTC Location: 42.33° N, 13.33° E, Depth: 8.8 km Magnitude: 5.8 Richter



The geodetic network

3 Italian IGS stations 32 stations in Abruzzo region 17 other stations within a distance of ~ 50 km from Abruzzo boundaries The geodetic network

3 Italian IGS stations 32 stations in Abruzzo region 17 other stations within a distance of ~ 50 km from Abruzzo boundaries

Data from 1st February (DOY 32) to 2nd, May (DOY 122) have been adjusted up to now from 32 to 95 (64 days): before earthquake from 96 to 122 (27 days): after earthquake

The geodetic network



ASI-Geodaf, INGV-RING, Leica-ItalPos, TopCon-Geotop, GPSAbruzzo, GPSUmbria, ResNap



The processing strategies 1/2

IGS stations stochastically constrained:

coordinates: interpolation of last 52 IGS05 weekly solutions,

> constraints: 2 mm horizontally, 4 mm in height

The processing strategies 1/2

IGS stations stochastically constrained:

coordinates: interpolation of last 52 IGS05 weekly solutions,

> constraints: 2 mm horizontally, 4 mm in height

Final IGS EOP, EPH and PCV's Adoption of the international standards in the raw data processing by BSW 5.0 software

The processing strategies 2/2



Outlier rejection

The processing strategies 2/2



Outlier rejection

Modelling the time series to estimate discontinuities

The processing strategies 2/2



Outlier rejection

Modelling the time series to estimate discontinuities

Spatial analysis of the discontinuities and clustering in subregions

Examples of time series: MATE



Examples of time series: MATE



Examples of time series: MEDI



Examples of time series: TERA



Examples of time series: TERA



Examples of time series: OCRA



Examples of time series: PAGA



Examples of time series: AQRA



Outliers rejection

Permanent networks are intrinsically redundant

to improve coordinates repeatabilities a severe automated outliers rejection is useful

Outliers rejection

Permanent networks are intrinsically redundant

to improve coordinates repeatabilities a severe automated outliers rejection is useful

This is a particular case: few data, manual analysis, ↓ conservative approach in outlier rejection just bad quality sessions before earthquake removed The results of IGS stations

3 stochastically constrained stations: CAGL, MATE, MEDI

Residuals of daily results wrt apriori coordinates							
(mm)	East	st North Heigh					
Mean	0.6	0.4	0.5				
σ	2.7	1.0	4.1				
Min	-4.0	-2.6	-10.8				
Max	6.4	6.6	17.4				

Time series interpretation (1/2)

Short time series in the geodetic analysis

constant model to avoid propagation of

seasonal effects and localized in time variations

into meaningless estimated velocities

Time series interpretation (2/2)

Before earthquake:

not a clear presence of pre seismic signal, just linear trend $$\downarrow\!\!\downarrow$$

linear trend estimation and removal

not to estimate velocities but to better model daily solutions

Time series interpretation (2/2)



After earthquake: a postseismic signal is often clear, but few days are available U at the present, simple costant model applied, with more data: linear and 2nd order polynomial

One example



(mm)	DE	DN	Dh
Constant	3.0	2.3	-2.9
Linear	0.9	0.1	-0.1

Residuals statistics of daily solutions

Before			After				
(mm)	Е	N	h	(mm)	Е	N	h
Mean	0.0	0.0	0.0	Mean	0.0	0.0	0.0
σ	1.5	1.3	3.6	σ	1.5	1.4	4.9
Min	-6.1	-6.3	-11.1	Min	-7.2	-6.5	-23.6
Max	7.1	5.9	11.4	Max	8.0	8.5	34.7

Worse height results after earthquake: post seismic assessment of 4 stations near L'Aquila **Parameters and covariances estimation**

Daily coordinates models in time ↓ Model parameters estimated by LS

Formal daily covariances typically underestimated and final covariances too much optimistic ↓ Empirical covariances estimation needed **Parameters and covariances estimation**

Few observations ↓ Simplified hypotheses on time series models and covariances Parameters and covariances estimation

Few observations ↓ Simplified hypotheses on time series models and covariances

Joint estimation of parameters and covariances U Typically an iterative process up to final results

$$\mathbf{y}_{0}, \tilde{\mathbf{C}}_{yy} \Longrightarrow \hat{\mathbf{x}}_{I}, \hat{\mathbf{C}}_{yy_{I}} \Longrightarrow \mathbf{y}_{0} \hat{\mathbf{C}}_{yy_{I}} \Longrightarrow \hat{\mathbf{x}}_{II}, \hat{\mathbf{C}}_{yy_{II}} \Longrightarrow ... \Longrightarrow \hat{\mathbf{x}}_{F}, \hat{\mathbf{C}}_{yy_{F}}$$

Hypotheses on network covariances

daily network covariance constant in time
 no correlations between consecutive days

$$\mathbf{C}(t_{k}) = \mathbf{C} \ \forall k = 1, ..., T, \underbrace{\mathbf{C}}_{3P \times 3P} = \begin{bmatrix} \mathbf{C}_{11} & \mathbf{C}_{12} & ... & \mathbf{C}_{1P} \\ \mathbf{C}_{21} & \mathbf{C}_{22} & ... & \mathbf{C}_{2P} \\ ... & ... & ... \\ \mathbf{C}_{P1} & \mathbf{C}_{P2} & ... & \mathbf{C}_{PP} \end{bmatrix},$$

$$\mathbf{C}_{ij} = \begin{bmatrix} c_{x_{1P_{i}}x_{1P_{j}}} & c_{x_{1P_{i}}x_{2P_{j}}} & c_{x_{1P_{i}}x_{3P_{j}}} \\ c_{x_{2P_{i}}x_{1P_{j}}} & c_{x_{2P_{i}}x_{2P_{j}}} & c_{x_{2P_{i}}x_{3P_{j}}} \\ c_{x_{3P_{i}}x_{1P_{j}}} & c_{x_{3P_{i}}x_{2P_{j}}} & c_{x_{3P_{i}}x_{3P_{j}}} \\ \end{bmatrix} = \begin{bmatrix} c_{ij} \\ ... \\ ... \\ \mathbf{C}_{ij} \\ ... \\$$

Estimation of the model parameters

Constant or linear model

$$\mathbf{X}_{P_i}(t) = \begin{cases} \mathbf{X}_{P_i}(\overline{t}) \\ \mathbf{X}_{P_i}(\overline{t}) + \dot{\mathbf{X}}_{P_i} \cdot (t - \overline{t}) \end{cases}$$

Estimation of the model parameters

Constant or linear model

$$\mathbf{x}_{P_i}(t) = \begin{cases} \mathbf{x}_{P_i}(\overline{t}) \\ \mathbf{x}_{P_i}(\overline{t}) + \dot{\mathbf{x}}_{P_i} \cdot (t - \overline{t}) \end{cases}$$

1. For each point i=1,...,P, each component I=1,2,3 an independent regression is estimated by Least Squares

$$\mathbf{y}_{0} = \begin{bmatrix} x_{l_{i0}}(t_{1}) \\ x_{l_{i0}}(t_{2}) \\ \dots \\ x_{l_{i0}}(t_{T}) \end{bmatrix}, \tilde{\mathbf{C}}_{yy} = \sigma_{0}^{2}\mathbf{I} \Longrightarrow \mathbf{LS} \Longrightarrow \hat{\mathbf{x}}_{I} = \begin{bmatrix} \hat{x}_{l_{i}}(\bar{t}) \\ \hat{x}_{l_{i}} \end{bmatrix}_{I}$$

Empirical estimation of the covariances

Estimated vector of the residuals

$$= \begin{bmatrix} x_{l_{i0}}(t_1) - [\hat{x}_{l_{iI}}(\overline{t}) + \hat{x}_{l_{iI}} \cdot (t_1 - \overline{t})] \\ x_{l_{i0}}(t_2) - [\hat{x}_{l_{iI}}(\overline{t}) + \hat{x}_{l_{iI}} \cdot (t_2 - \overline{t})] \\ \dots \\ x_{l_{i0}}(t_T) - [\hat{x}_{l_{iI}}(\overline{t}) + \hat{x}_{l_{iI}} \cdot (t_T - \overline{t})] \end{bmatrix}$$

Empirical estimation of the covariances

Estimated vector of the residuals

$$f_{l_{iI}} = \begin{bmatrix} x_{l_{i0}}(t_1) - [\hat{x}_{l_{iI}}(\overline{t}) + \hat{x}_{l_{iI}} \cdot (t_1 - \overline{t})] \\ x_{l_{i0}}(t_2) - [\hat{x}_{l_{iI}}(\overline{t}) + \hat{x}_{l_{iI}} \cdot (t_2 - \overline{t})] \\ \dots \\ x_{l_{i0}}(t_T) - [\hat{x}_{l_{iI}}(\overline{t}) + \hat{x}_{l_{iI}} \cdot (t_T - \overline{t})] \end{bmatrix}$$

Estimated covariances and correlations

$$\hat{c}_{ij_{lm I}} = \frac{1}{T - N} \hat{\mathbf{r}}_{i_{I}}^{T} \hat{\mathbf{r}}_{m_{j_{I}}}$$

(N=1/2 for the constant/ linear model)

Final results

With the above hypotheses, no need of iterations



Final results

With the above hypotheses, no need of iterations



Propagation of coordinates and covariances

Displacement at earthquake epoch

$$\hat{\mathbf{x}}_{iB}(t_E) = \hat{\mathbf{x}}_i(\overline{t_B}) + \hat{\dot{\mathbf{x}}}_i(t_E - \overline{t_B})$$
$$\hat{\mathbf{x}}_{iA}(t_E) = \hat{\mathbf{x}}_{iA}(\overline{t_A})$$
$$\delta \hat{\mathbf{x}}_i(t_E) = \hat{\mathbf{x}}_{iA}(t_E) - \hat{\mathbf{x}}_{iB}(t_E)$$

Propagation of coordinates and covariances

Displacement at earthquake epoch

$$\hat{\mathbf{x}}_{iB}(t_E) = \hat{\mathbf{x}}_i(\overline{t_B}) + \hat{\dot{\mathbf{x}}}_i(t_E - \overline{t_B})$$
$$\hat{\mathbf{x}}_{iA}(t_E) = \hat{\mathbf{x}}_{iA}(\overline{t_A})$$
$$\delta \hat{\mathbf{x}}_i(t_E) = \hat{\mathbf{x}}_{iA}(t_E) - \hat{\mathbf{x}}_{iB}(t_E)$$

Covariance of the displacement $\mathbf{C}_{iB}(t_E) = \mathbf{C}_{\overline{xx}iB} + \mathbf{C}_{\dot{x}\dot{x}i}(t_E - \overline{t}_B)^2$ $\mathbf{C}_{iA}(t_E) = \mathbf{C}_{\overline{xx}iA}$ $\mathbf{C}_{i\delta\delta}(t_E) = \mathbf{C}_{iA}(t_E) + \mathbf{C}_{iB}(t_E)$

Covariances of the two propagations







Separation of rigid motion from deformation



BAD SPATIAL INTERPOLATION

Separation of rigid motion from deformation



BAD SPATIAL INTERPOLATION



GOOD SPATIAL INTERPOLATION

PIECEWISE INTERPOLATION INVOLVES DISCONTINUITIES = FAULTS !

Spatial covariances and interpolation



A signal could be isolated, but quite arbitrarily

a preliminary clustering of homogeneous areas needed

L'Aquila sites: 20-70 mm W displacements





L'Aquila sites:
 20-70 mm W
 displacements

2. Around them: smaller S-W displacements



L'Aquila sites:
 20-70 mm W
 displacements

2. Around them: smaller S-W displacements

East region:
 2-30 mm NE
 displacements



1. L'Aquila sites : 20-70 mm W displacements

2. Around them: smaller S-W displacements

3. East region:2-30 mm NEdisplacements

4. Around it:
no significant displacements,
but consistent directions.



1. L'Aquila sites : 20-70 mm W displacements

2. Around them: smaller S-W displacements

3. East region:2-30 mm NEdisplacements

4. Around it:
no significant displacements,
but consistent directions.

5. No other significant displacements

Separation of rigid motion from deformation

Horizontal motion of a network on earth surface: rotation of all the points around an axis with angular velocity ω



Separation of rigid motion from deformation

Horizontal motion of a network on earth surface: rotation of all the points around an axis with angular velocity ω



$$\mathbf{v}_i = [\boldsymbol{\omega} \times] \mathbf{x}_i$$

ω can be estimated
 by minimization of relative kinetic
 energy of the network

$$T_{\rm ap} = \sum_{i=1,\dots,P} \mathbf{v}_i^{\rm T} \mathbf{v}_i = \min$$

Realization of a Discrete Tisserand reference system Horizontal analysis in separate regions

Probably no significant rotation of networks but differential displacements

up to now no Tisserand analysis but statistics on displacements for the two main regions

East (14 stations)			L'Aquila (4 stations)				
(mm)	E	Ν	2D	(mm)	Е	Ν	2D
Mean	7.0	7.6	10.7	Mean	-40	1	41
σ	4.5	6.6	7.5	σ	22	8	21
Min	1.0	2.1	3.4	Min	-66	-11	19
Max	14.4	27.5	30.1	Max	-16	8.5	66

Vertical displacements





Conclusions

6th April earthquake in L'Aquila has been accompanied by an extension along an axis oriented NW-SE:

L'Aquila area and an Eastern Adriatic area interested by significant opposite horizontal displacements

Conclusions

6th April earthquake in L'Aquila has been accompanied by an extension along an axis oriented NW-SE:

L'Aquila area and an Eastern Adriatic area interested by significant opposite horizontal displacements

Significant gradients in the horizontal displacements of the Eastern Adriatic area

L'Aquila sites interested by vertical displacements of about 10 cm

Future analyses

Longer time series, to: increase the populations after the earthquake, analyze the post seismic time series

Future analyses

Longer time series, to: increase the populations after the earthquake, analyze the post seismic time series

More rigorous clustering in separate regions, rigorous Tisserand analysis, geometric analysis in Adriatic region

Future analyses

Longer time series, to: increase the populations after the earthquake, analyze the post seismic time series

More rigorous clustering in separate regions, rigorous Tisserand analysis, geometric analysis in Adriatic region

Cross comparison in L'Aquila with SAR interferograms